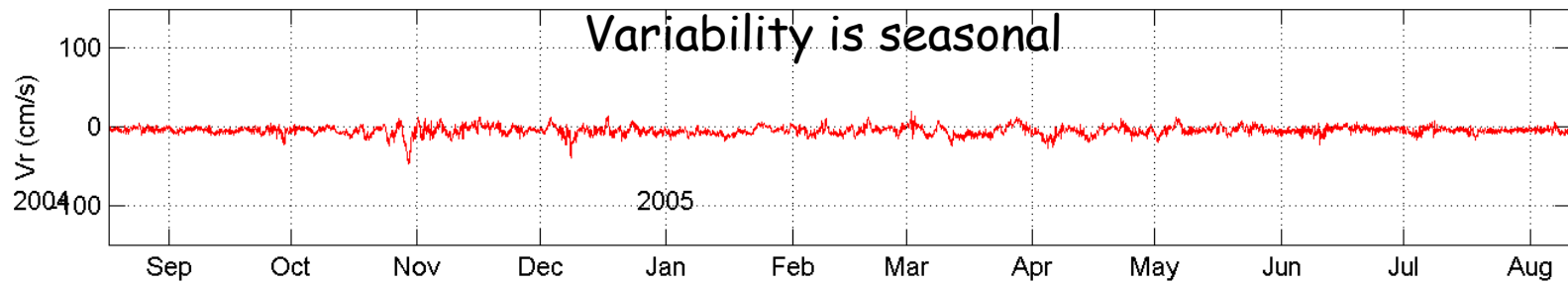
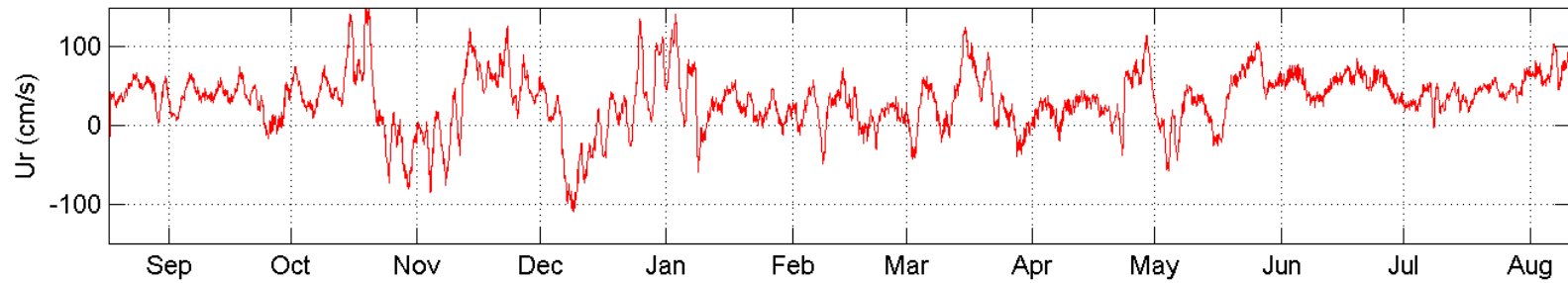
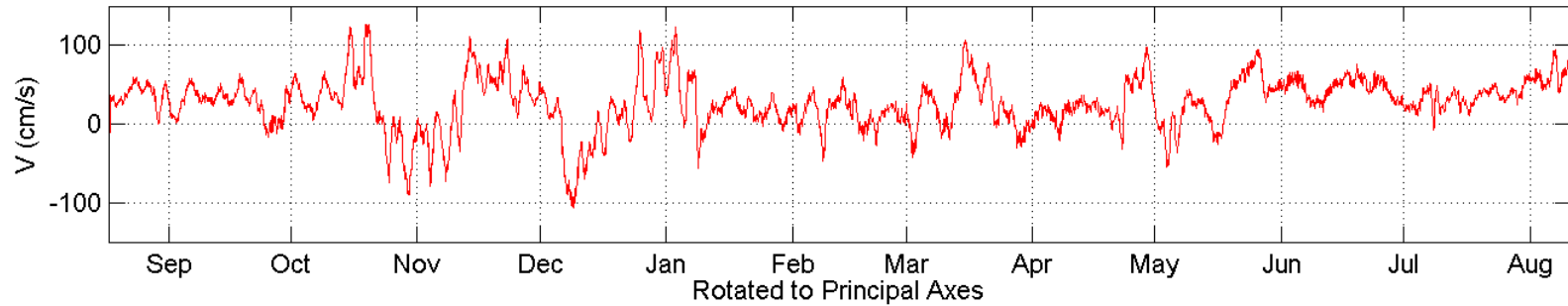
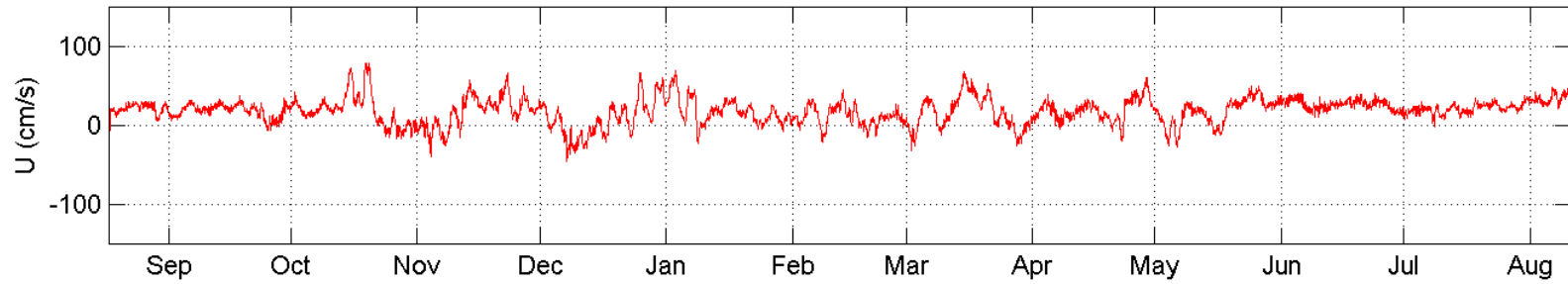
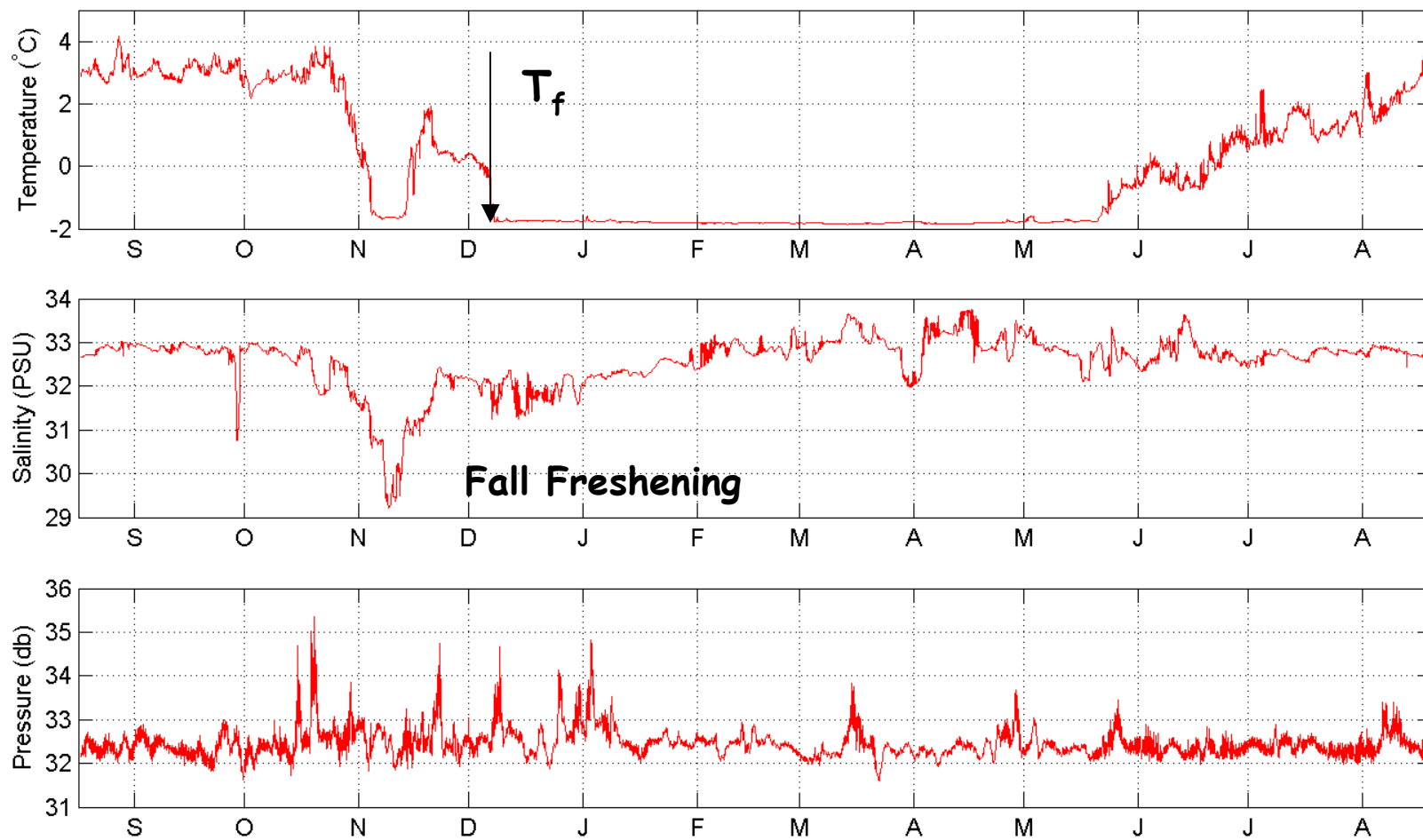


Western Bering Strait RCM9, 2004-2005 Preliminary Data

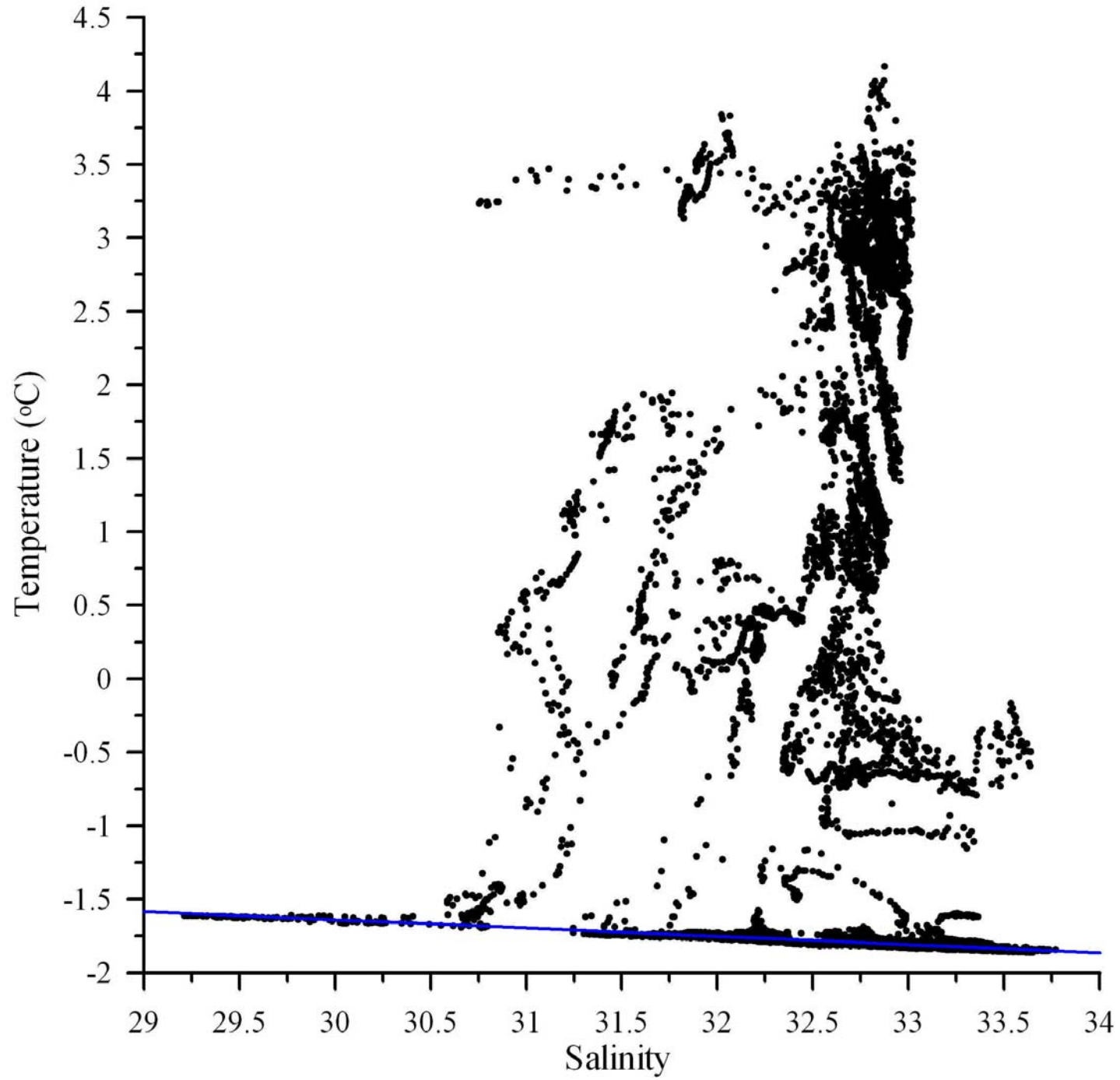


Western Bering Strait Mooring, 2004-2005 Preliminary Data

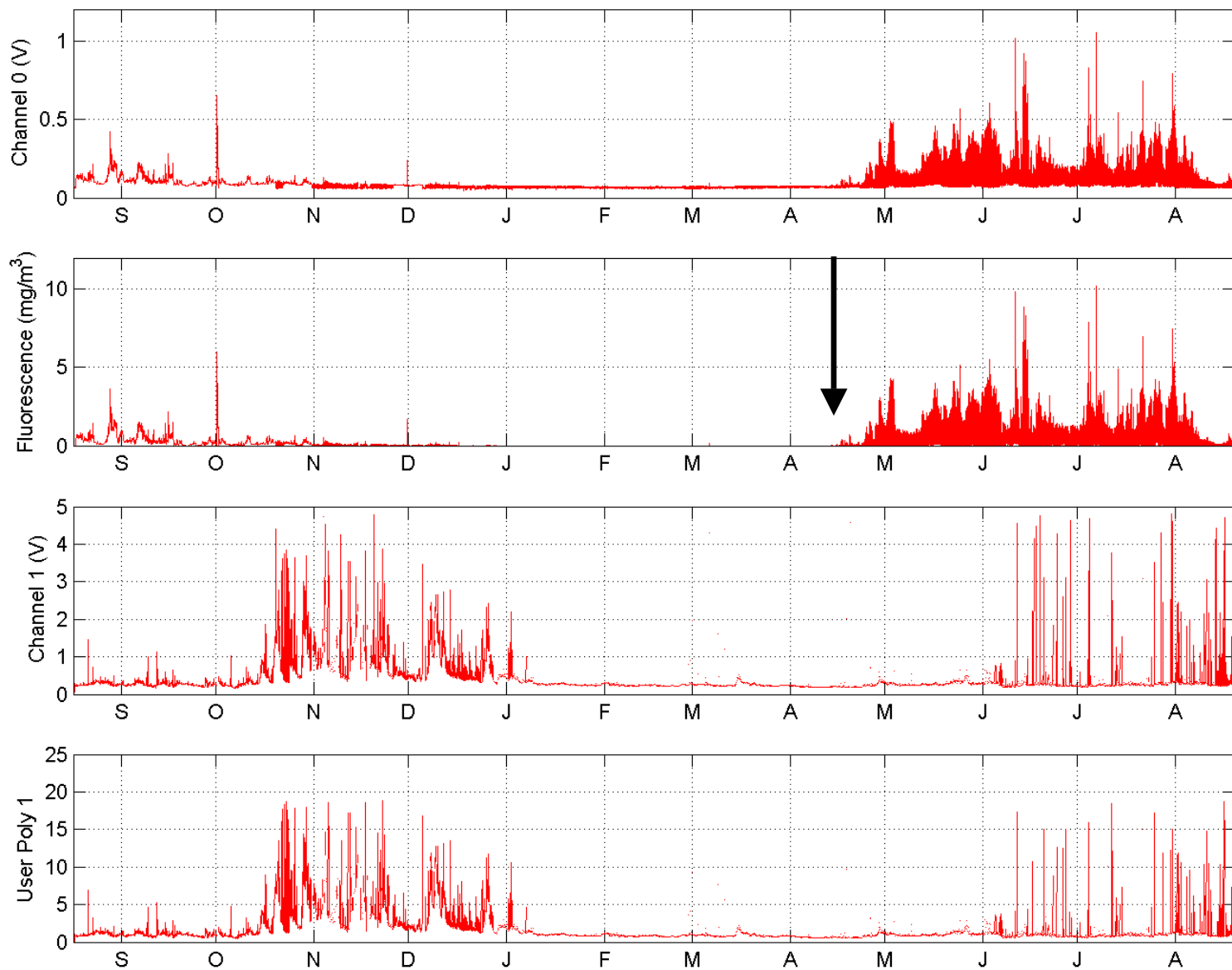


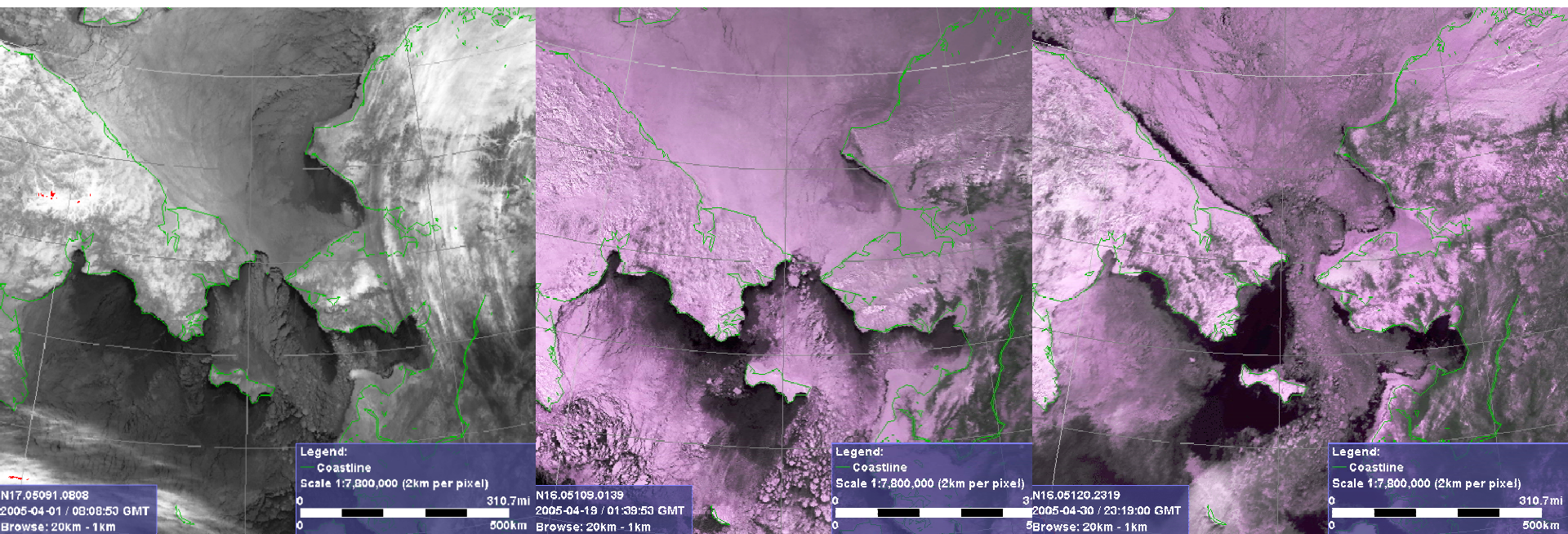
2004

2005



Western Bering Strait Mooring, 2004-2005 Preliminary Data



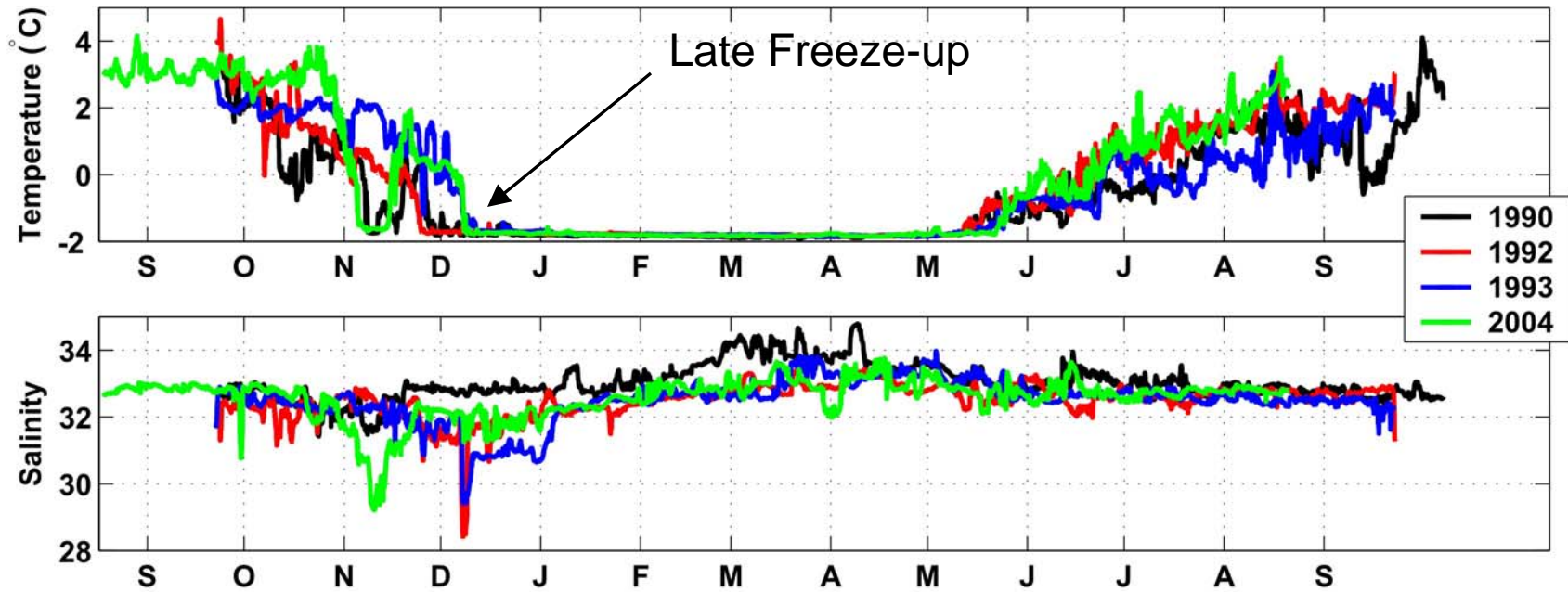


April 1

April 19  
2005

April 30

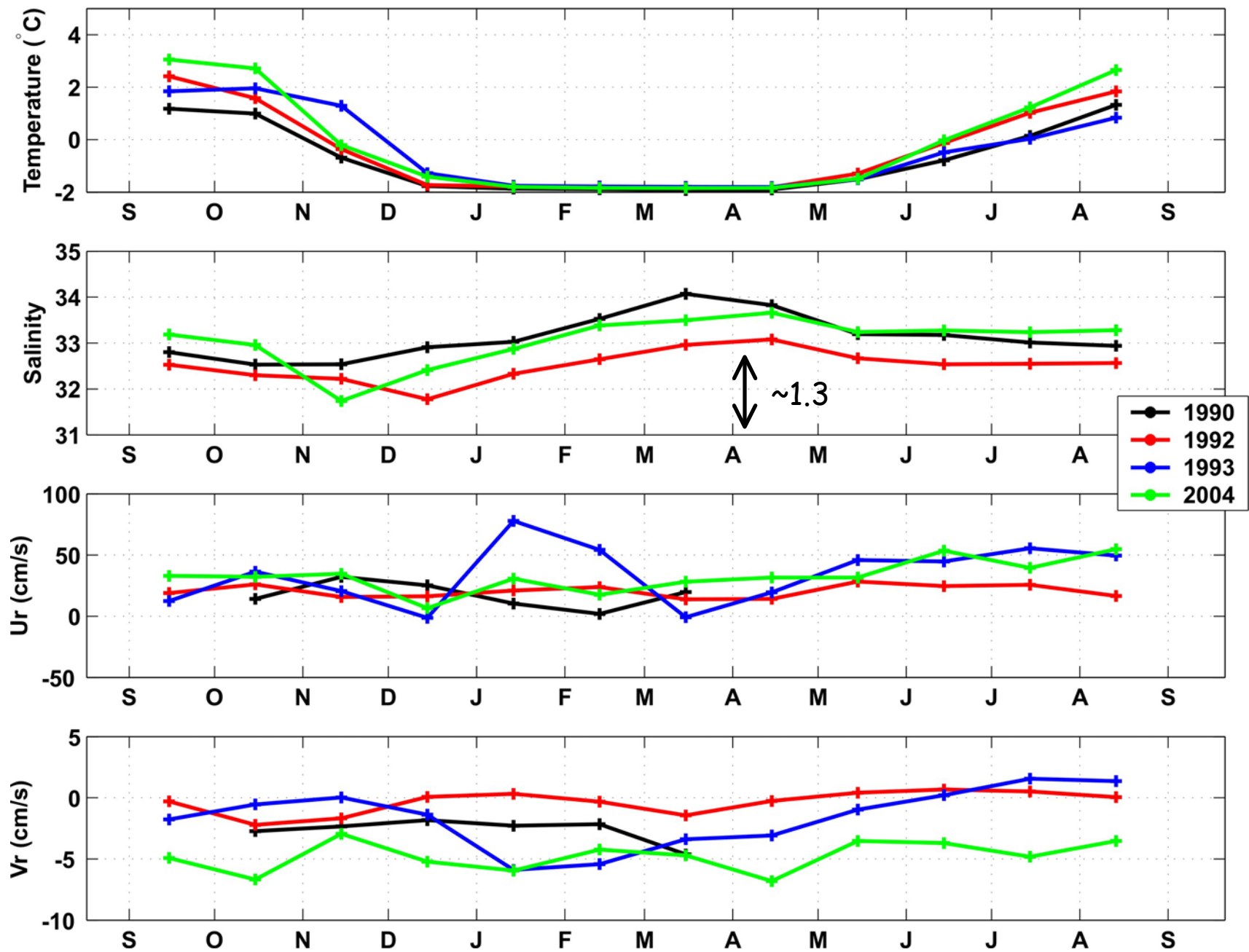
Mooring A1, SBE

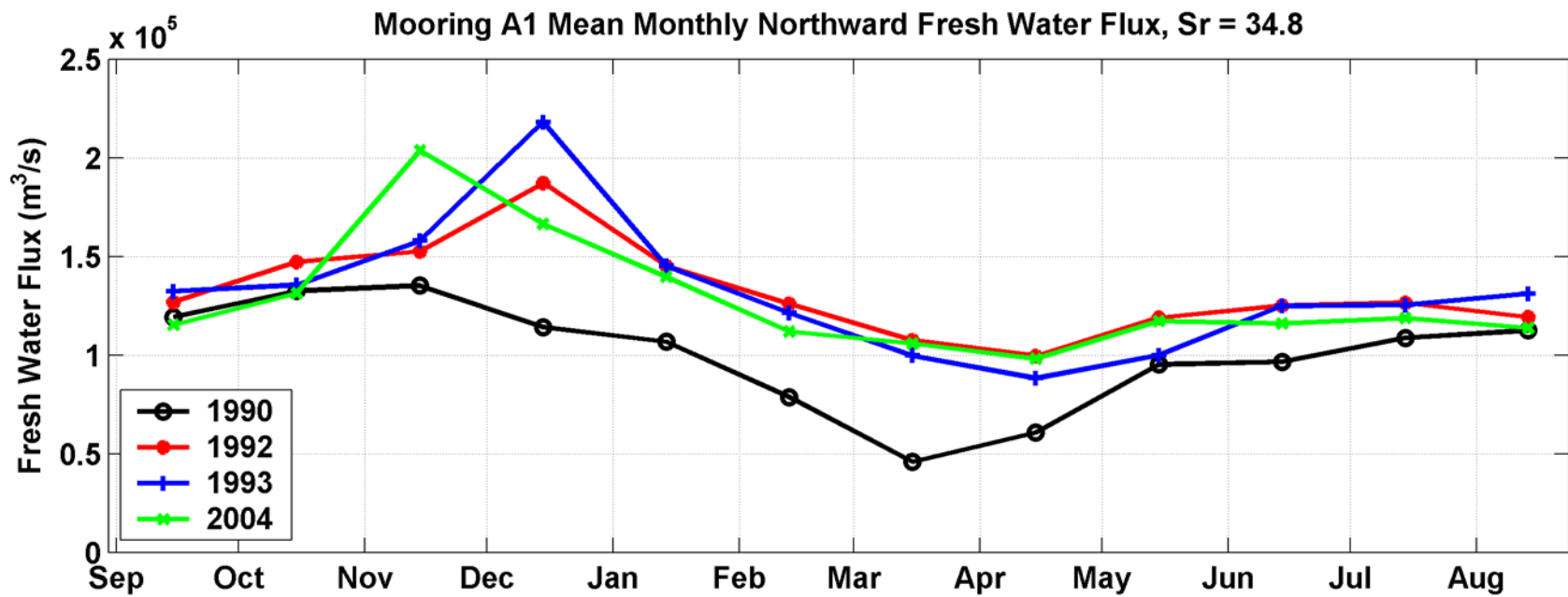
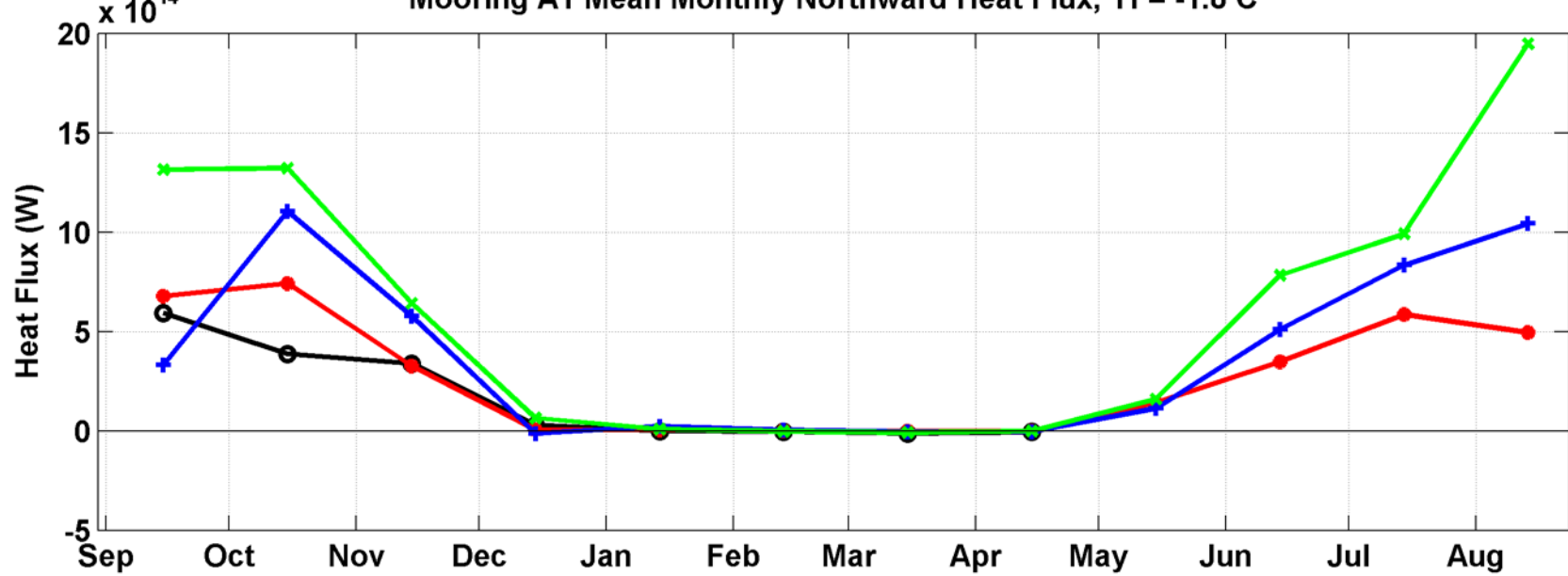


**Fall Freshening  
events  
Except 1990**



# Mooring A1, RCM





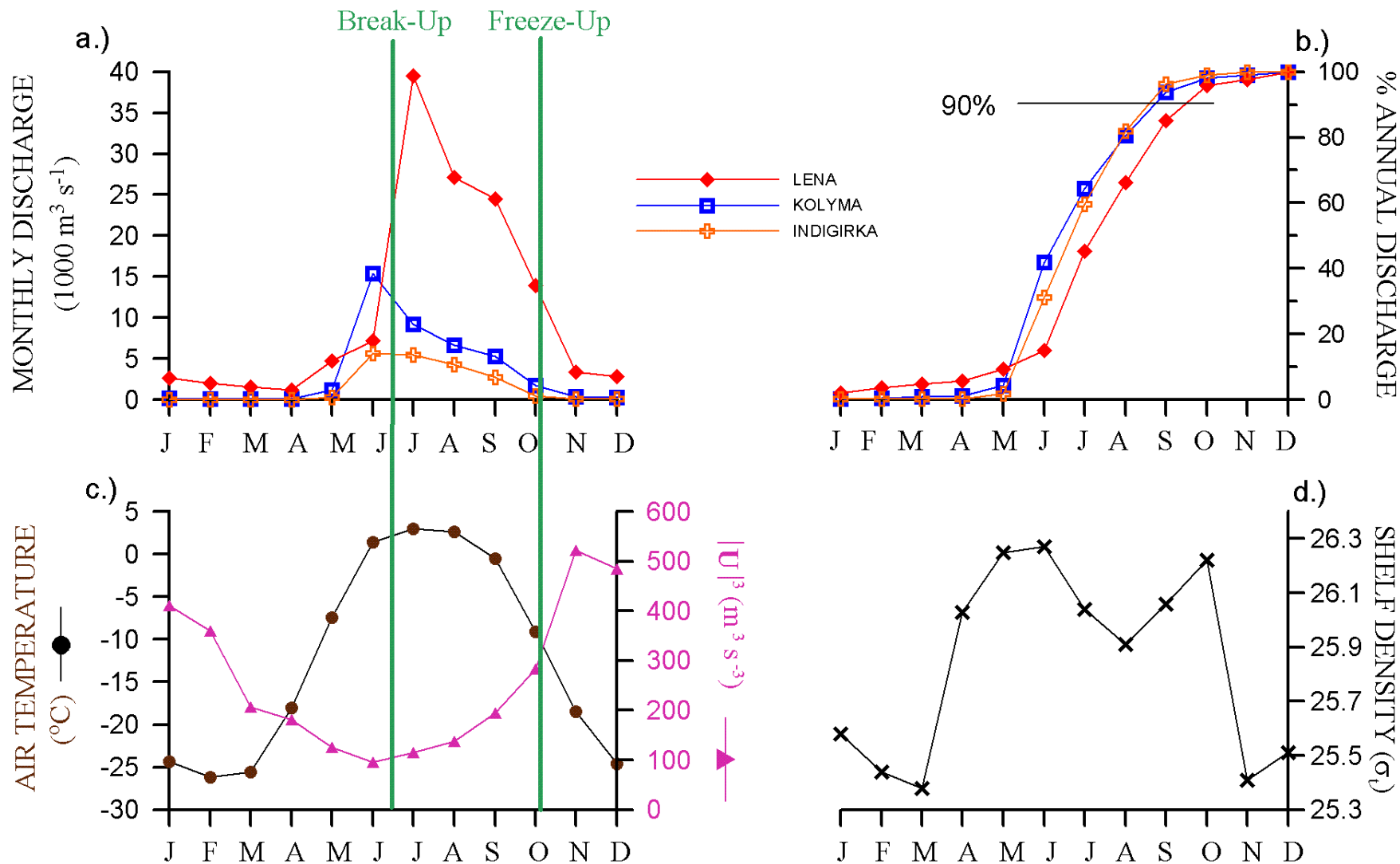


## Comments:

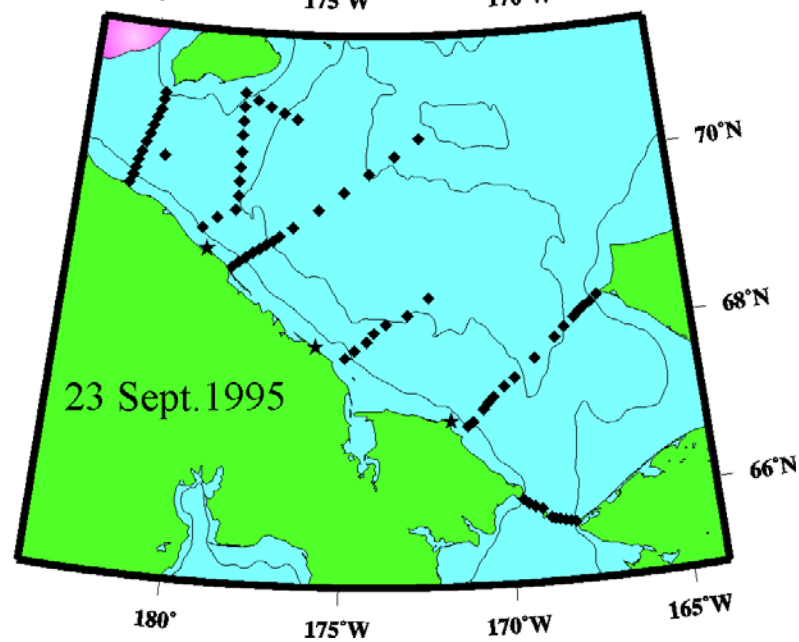
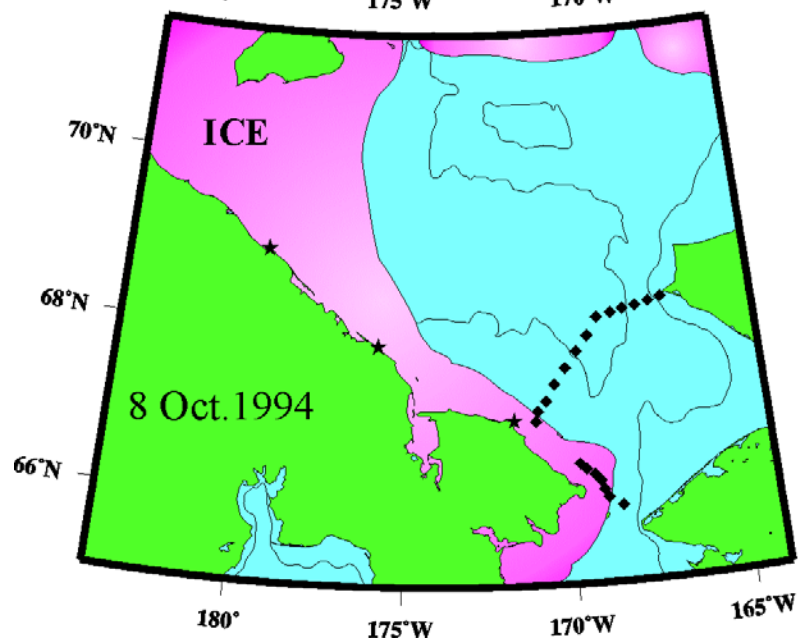
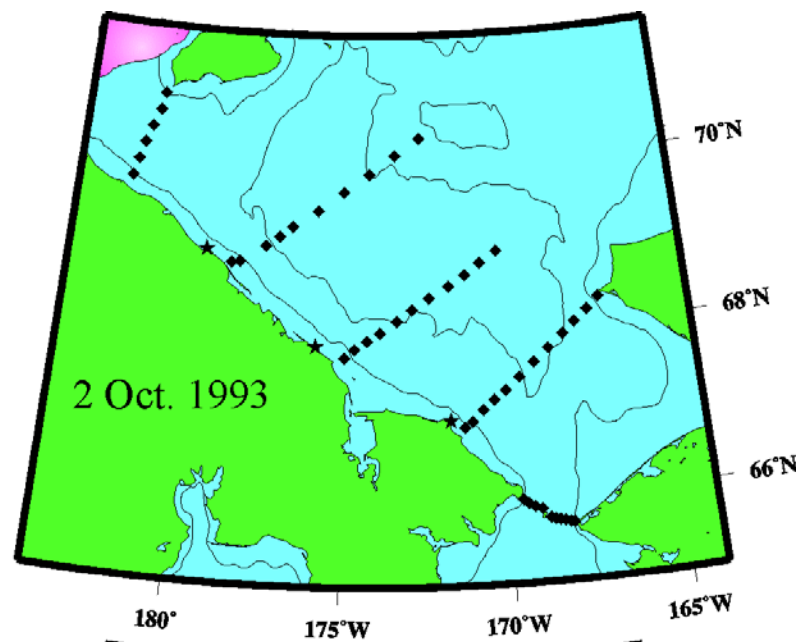
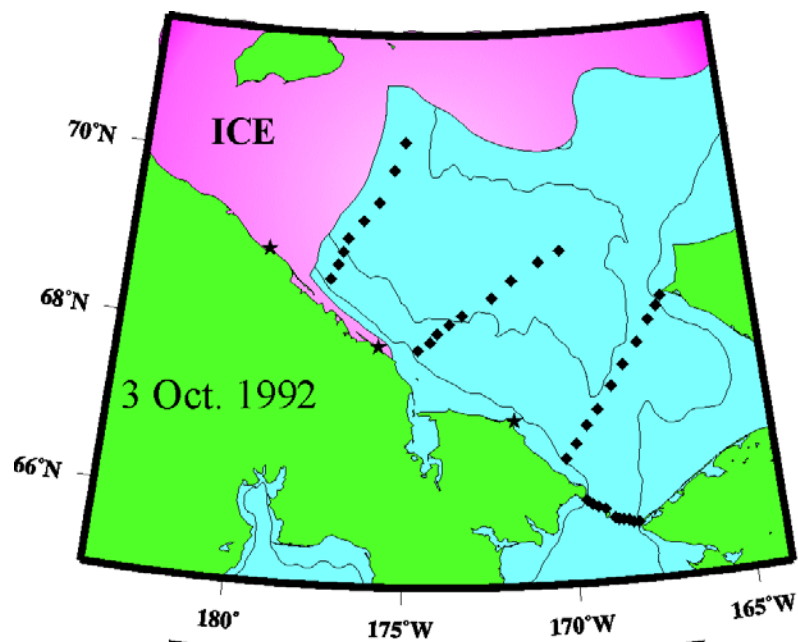
1. Data looks good
2. Temperatures in late summer-fall 2004 were  $\sim 1^\circ\text{C}$  warmer than previous years. Salinity not anomalous
3. Northward heat fluxes 2-3 times greater - affects onset of fall freeze-up and heat export into the Arctic Ocean.
4. Fall freshening events appear to affect the winter salinity increase (and thus halocline ventilation). Freshening is a 3-D process and not limited to the strait.
5. Fall northward flux of marine snow? Carbon transport (type and quantity) varies seasonally.
6. Fluorescence begins in mid-April - is it tied to leads?
7. Why is the duration of the waters with  $T = T_f$  in the strait constant (Dec-May) although the Bering Sea ice edge extent has varied tremendously?



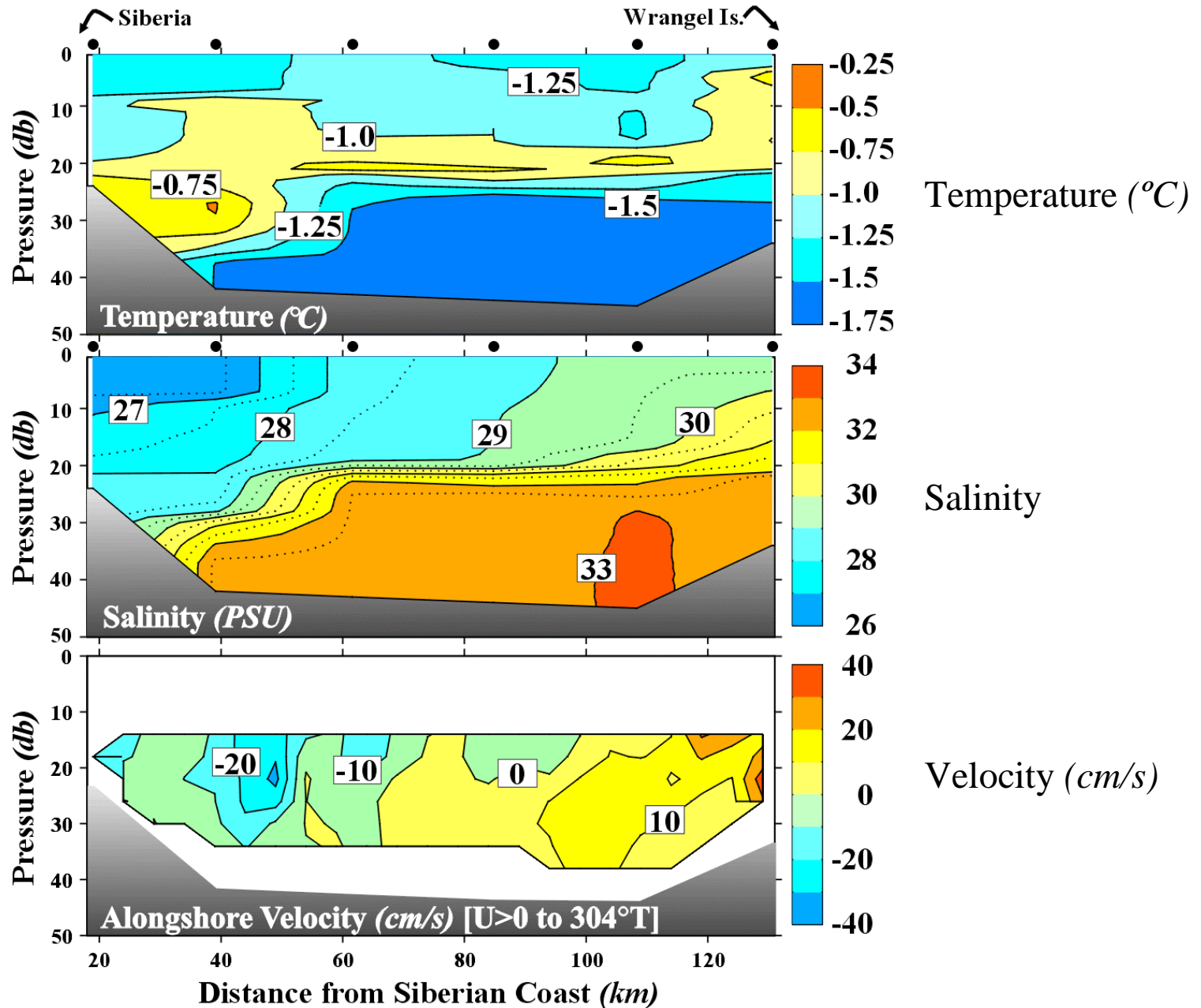
# ANNUAL CYCLE OF CLIMATOLOGICAL FACTORS THAT AFFECT MIXING



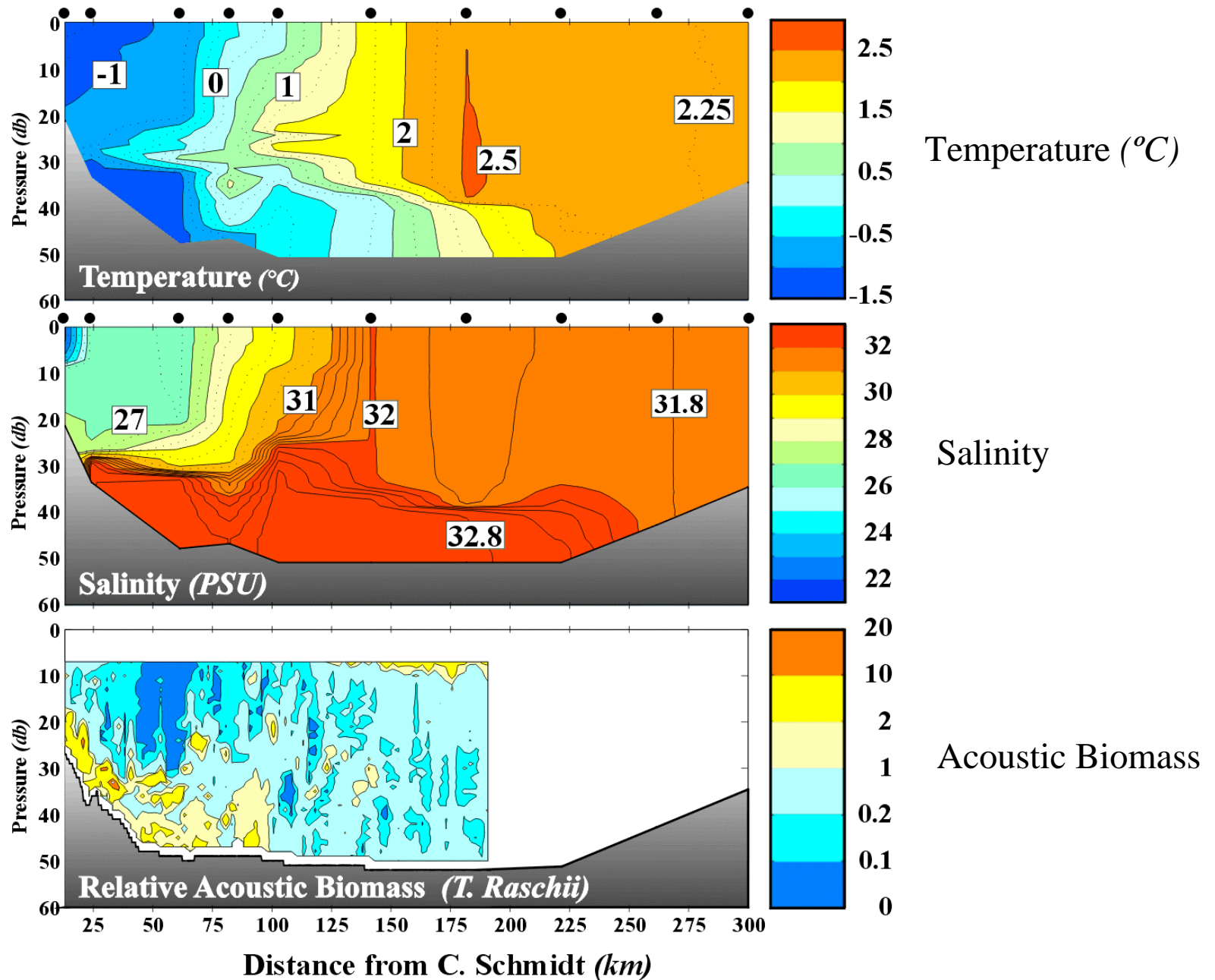
a) Mean monthly discharge, b) cumulative percentage discharge, c) air temperature and  $|U|^3$  (proportional to the rate of working by the wind on the water), and d) the annual cycle of shelf water density (as measured at a site on the central Chukchi shelf).



# Long Strait: October 1993

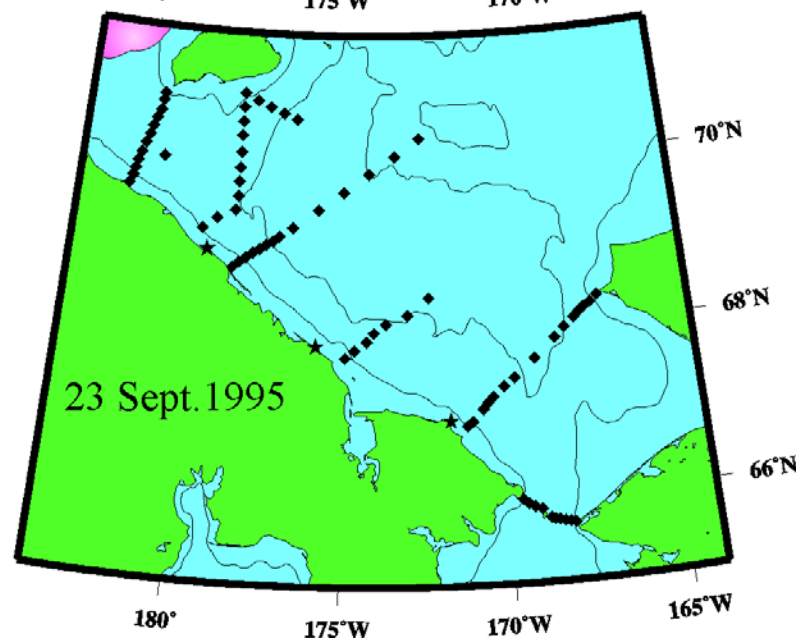
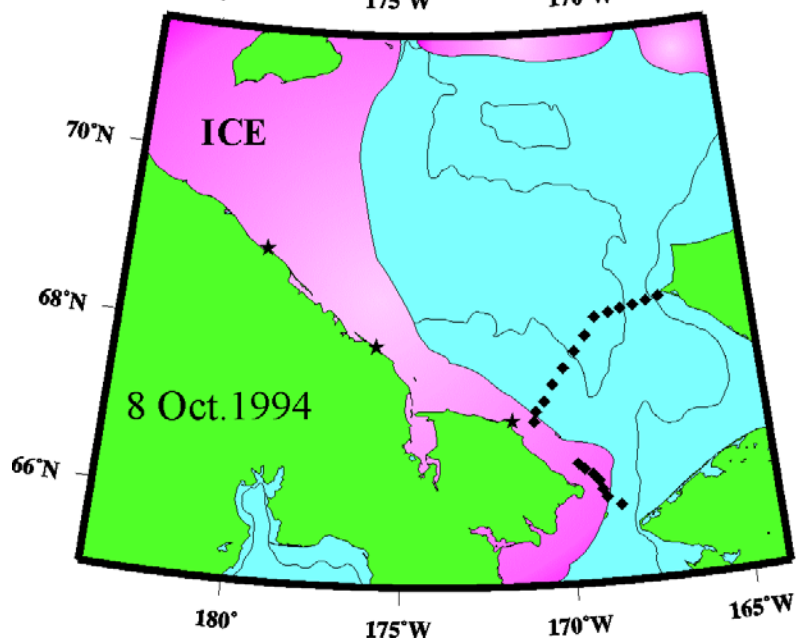
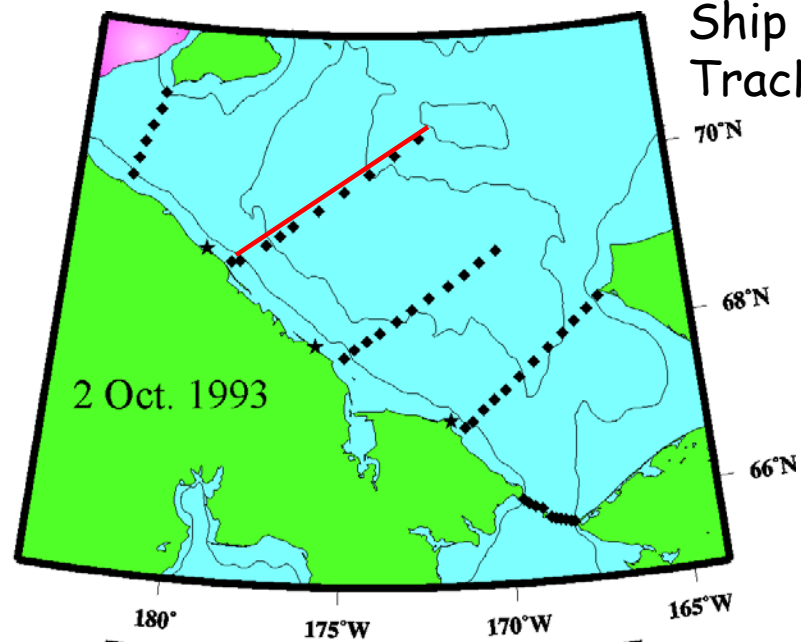
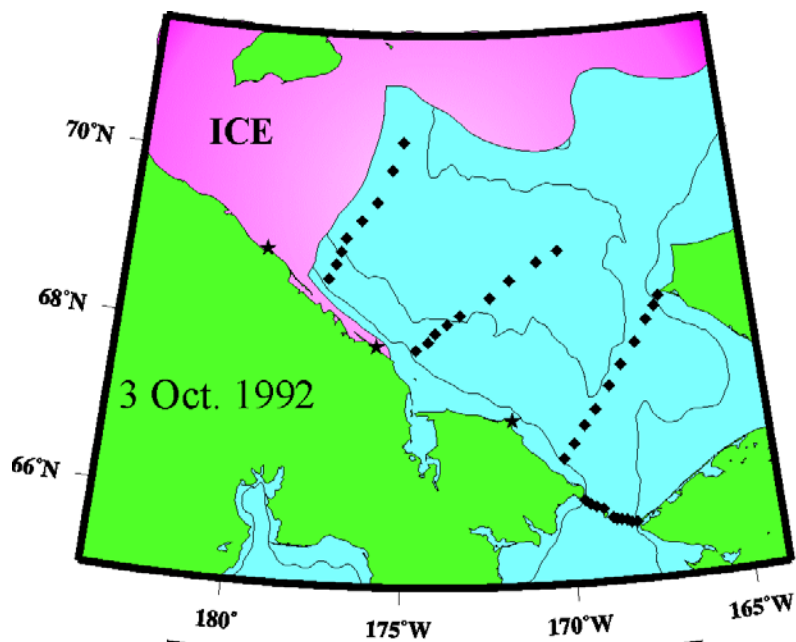


October 1993

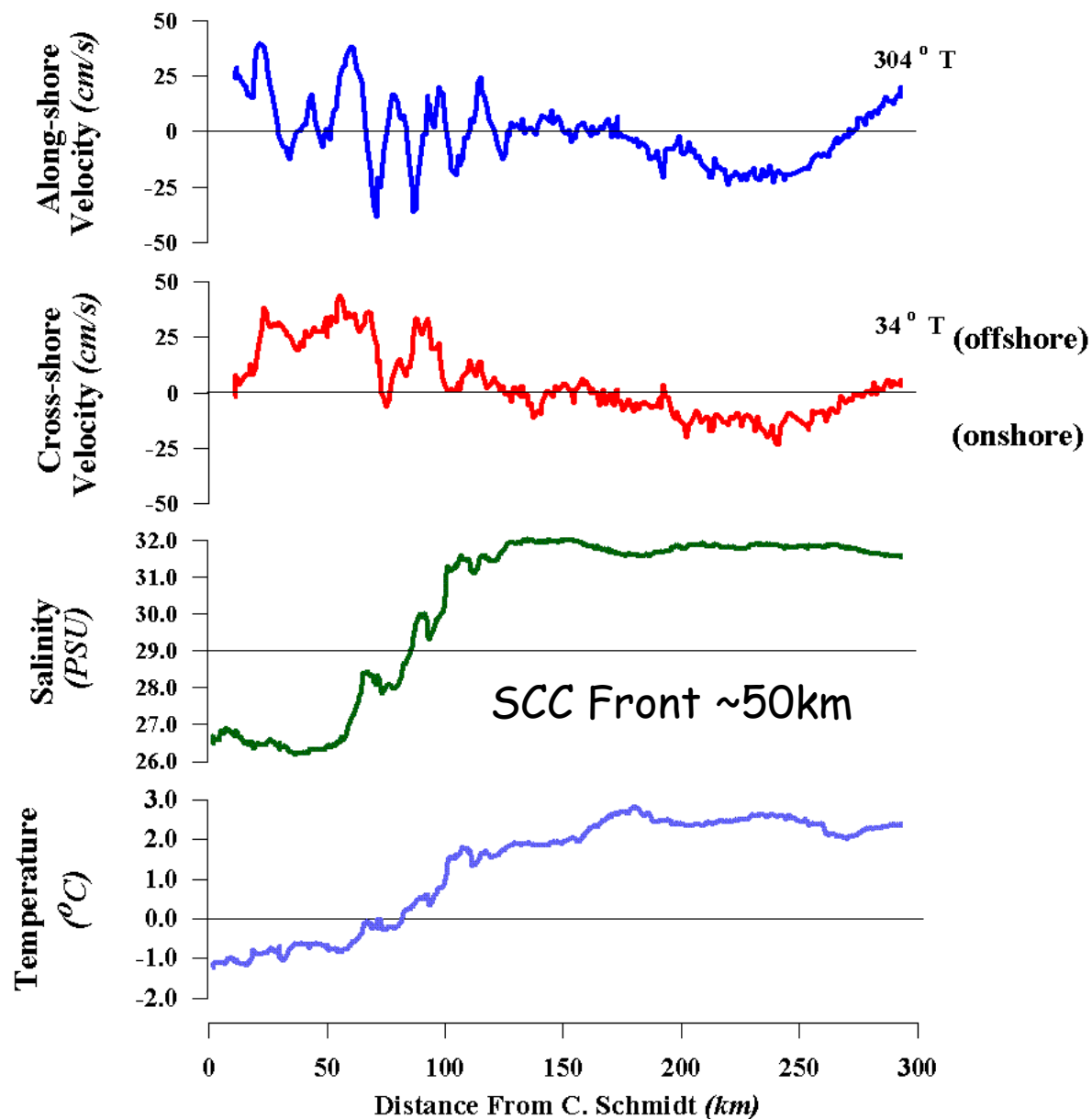


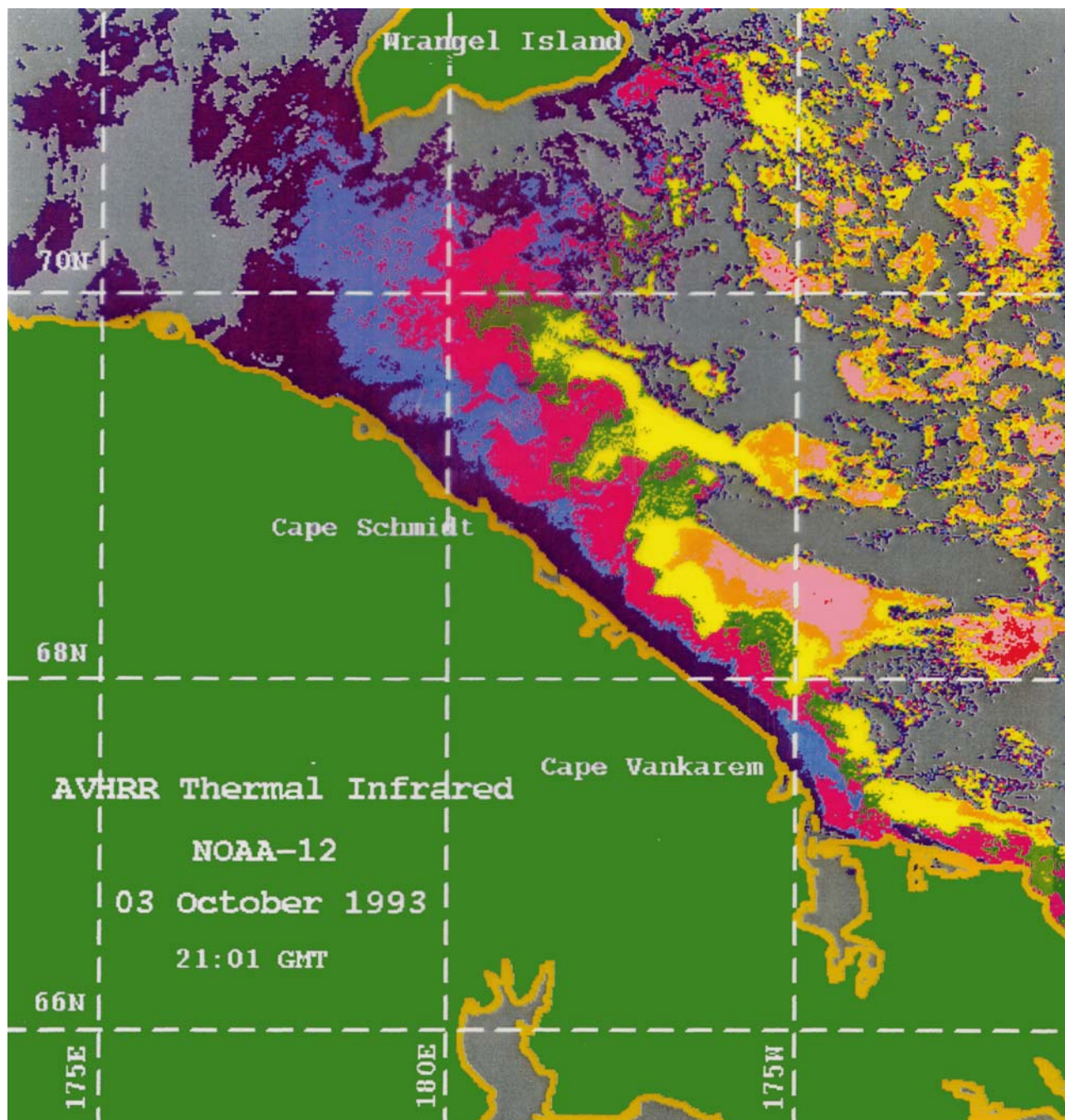


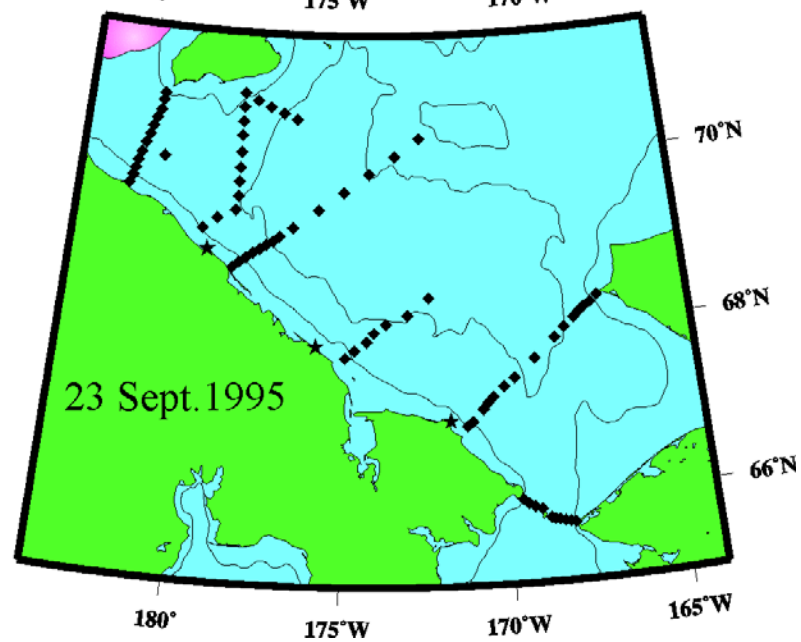
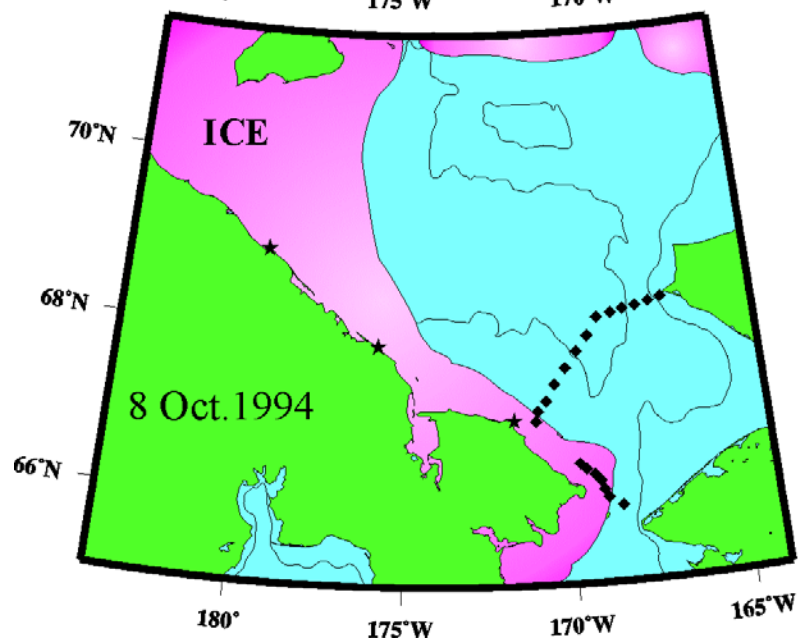
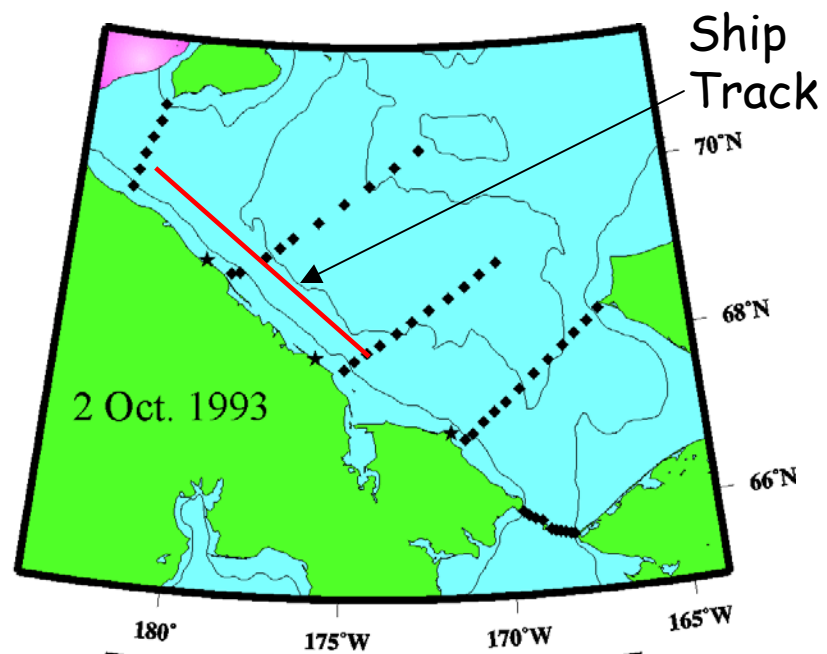
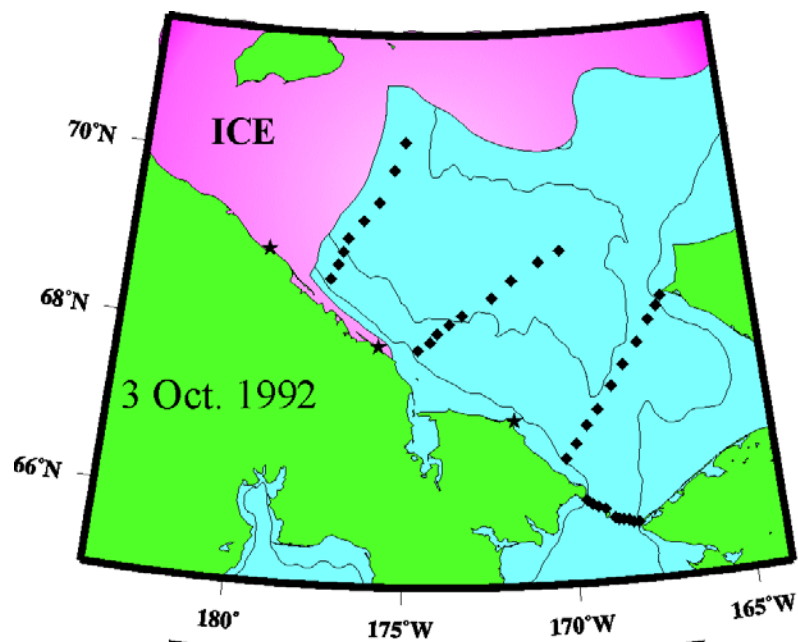
# Ship Track



# October 1993



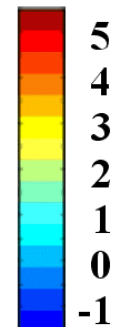
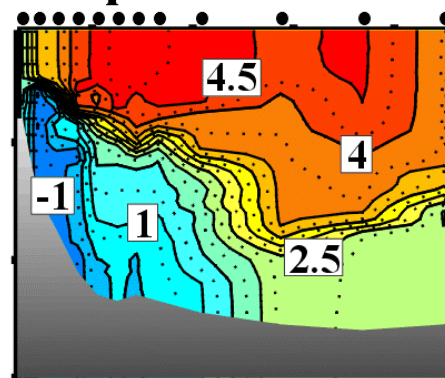
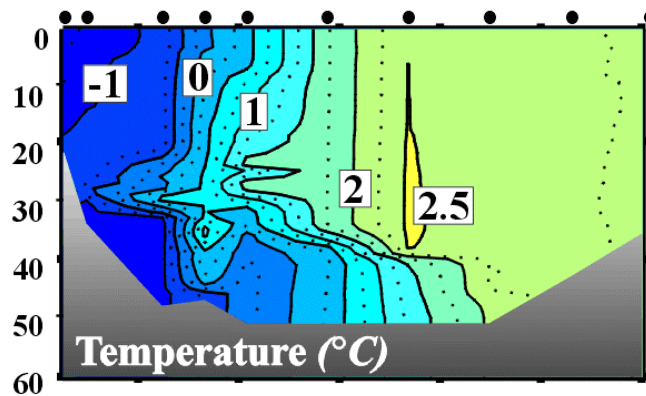




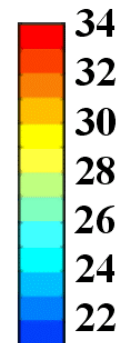
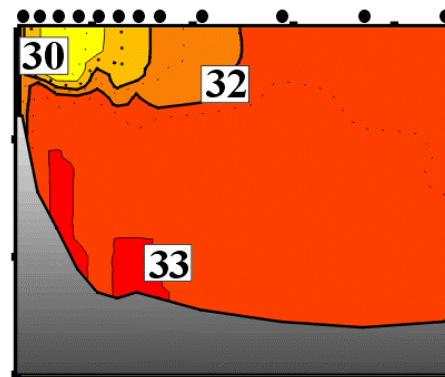
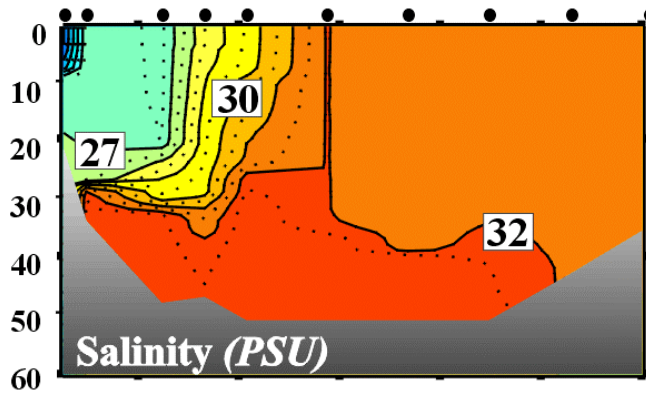


October 1993

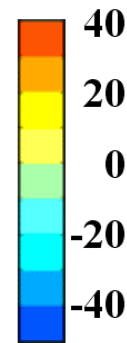
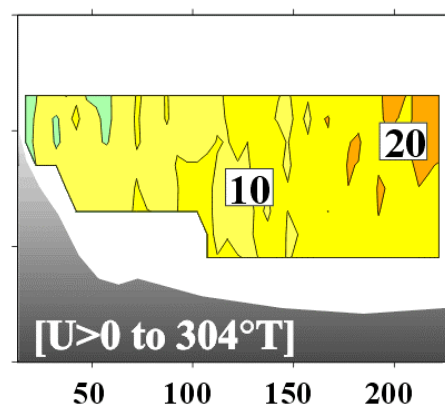
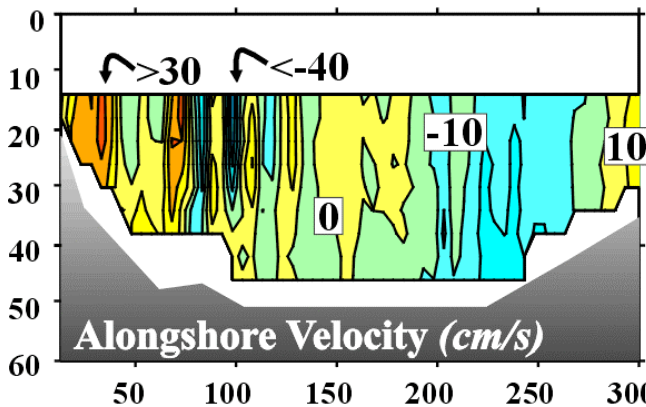
September 1995



Temperature ( $^{\circ}\text{C}$ )



Salinity



Velocity ( $\text{cm/s}$ )

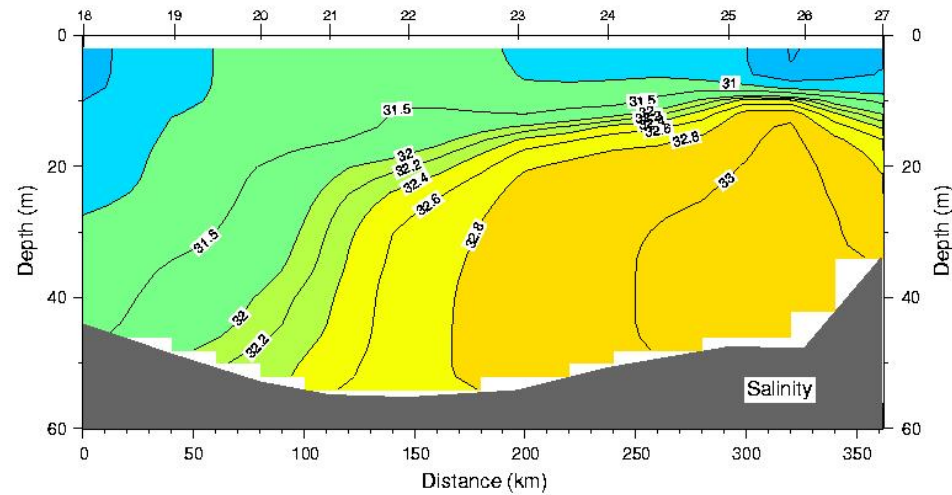
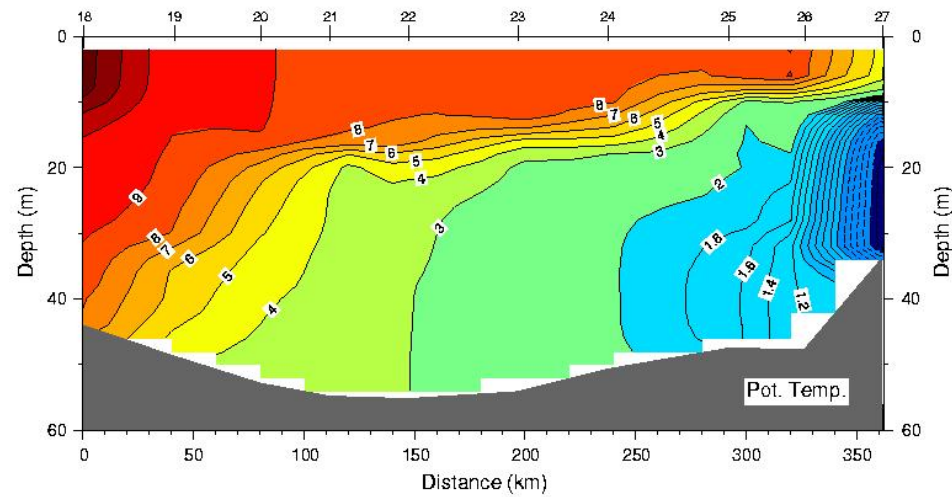
Distance from C. Schmidt (km)

# RUSALCA 2004

Properties [Section 3f]

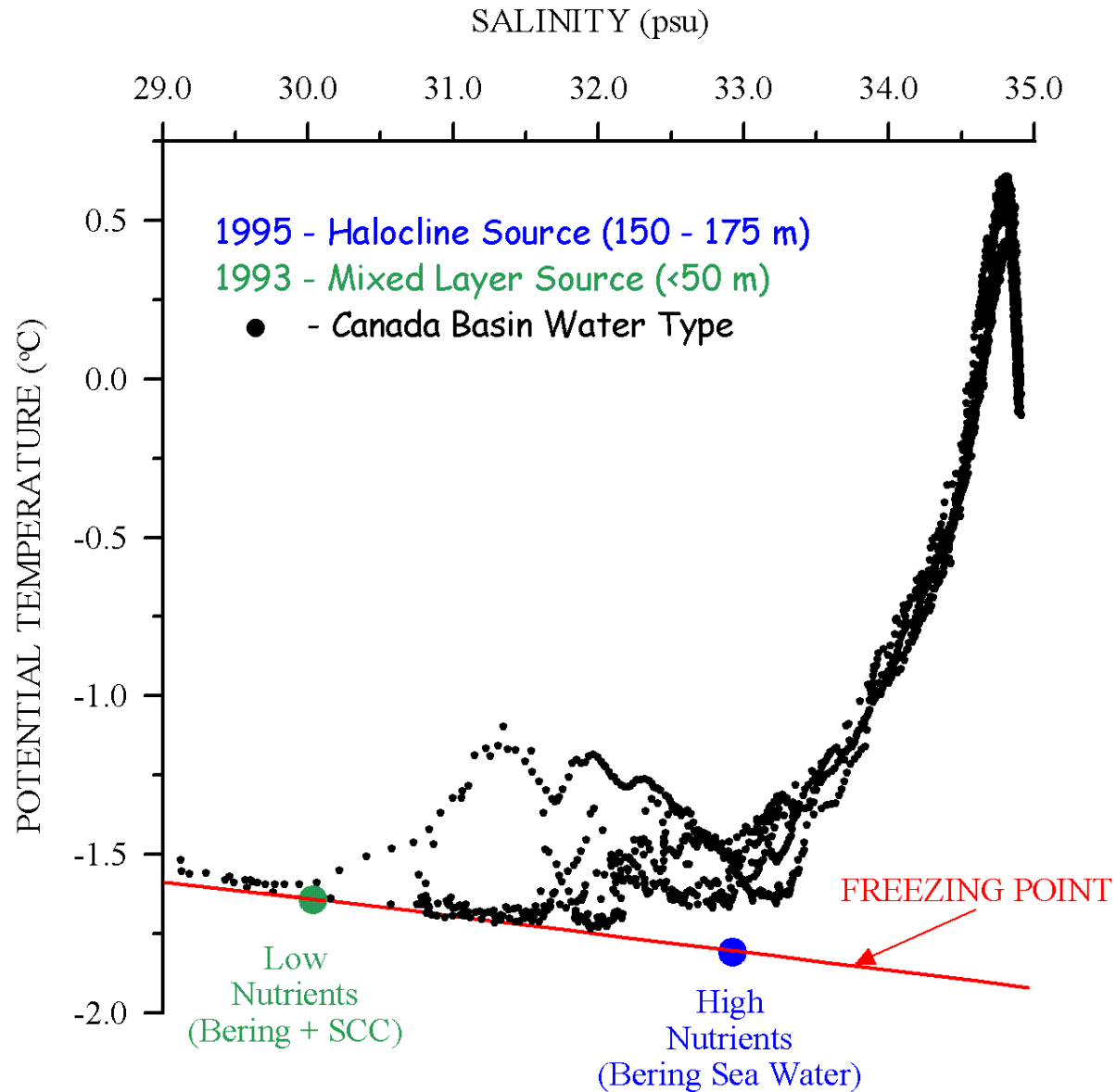
Alaska

Siberia





HYPOTHETICAL WINTER WATER MASSES  
WESTERN CHUKCHI SEA [1993](#) VS. [1995](#)

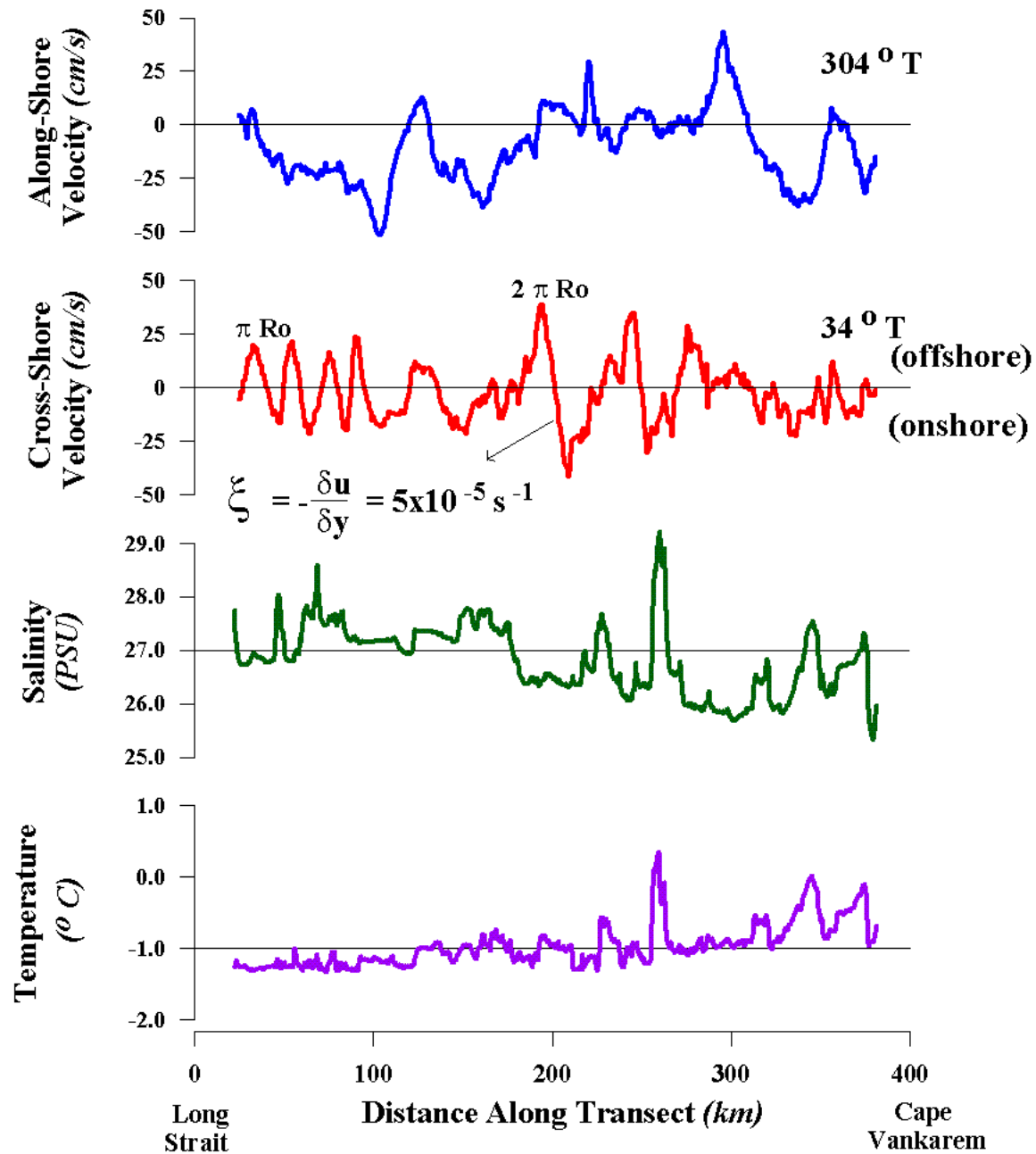


The SCC is a broad current carrying Siberian river water southwestward onto the Chukchi shelf. It has a well-developed front that is unstable (energetic eddies and meanders) . Habitats? Cross-shore transports?

It does not appear to have been well-developed during RUSLACA 2004 - the unusual situation. ( Bering Sea Water instead of Siberian Coastal Water along Chukotka).

So how representative is RUSALCA 2004 with respect to the near-shore environment?

# October 1993



## EDDY BUOYANCY FLUXES

Long Strait - Cape VanKarem :  $\overline{u' \rho'} = -2 \times 10^{-2} \text{ kg-m}^{-2}$

$$\iint \overline{u' \rho'} dz dy = -1.5 \times 10^5 \text{ kg-s}^{-1} \text{ out of SCC}$$

$$dz = 20 \text{ m}; dy = 400 \text{ km}$$

Jul - Sept. average discharge (Kolyma +Indigirka):  $Q = 15000 \text{ m}^3\text{-s}^{-1}$

$$\rho' \sim -20 \text{ kg-m}^{-3}$$

$$Q\rho' \sim 3 \times 10^5 \text{ kg-s}^{-1} \text{ to SCC}$$

Moreover, from *Spall and Chapman* [1998]:

$$\overline{u' \rho'} = c_e V_m \Delta \rho; \quad c_e \sim 0.02 - 0.04$$

Long Strait:  $V_m : 20 \text{ cm-s}^{-1}$ ;  $\Delta \rho : 4 \text{ kg-m}^{-3}$

Long Strait - Cape VanKarem :  $\overline{u' \rho'} = -2 \times 10^{-2} \text{ kg-m}^{-2}$

$$c_e \sim 0.02 - 0.03$$

Suggesting:

1. The instabilities might be important for the offshore dispersal of freshwater, and
2. The Spall and Chapman parameterization might usefully be applied in Arctic GCMs.