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PMEL Tsunami Forecast Series: Vol. 68 A Tsunami Forecast Model for Virginia Beach, Virginia

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PMEL Tsunami Forecast Series: Vol. 68 A Tsunami Forecast Model for Virginia Beach, Virginia

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5 Abstract

3 4

6 This report documents the development and testing of a tsunami forecast model for Virginia Beach, Virginia as part of NOAA's operational tsunami forecast system. Based 7 on the Method of Splitting Tsunamis (MOST) model, the forecast model performs 8 9 calculations on bathymetry grids with a horizontal resolution of 2 arc-seconds. At this 10 resolution it is capable of simulating four hours of tsunami wave dynamics in under ten 11 minutes of computational time. A reference inundation model of higher resolution (1/3 12 arc-seconds) was also developed in parallel, to provide a reference for the forecast model. 13 Both models were tested with nine simulated magnitude-9.3 tsunamis from different 14 source regions and one micro tsunami. Good agreement was observed between the 15 model computations, and the numerical consistency between the model results for the 16 maximum amplitudes and currents indicate reliability, robustness, and stability of the 17 forecast model.

18

The study shows that mega tsunamis from subduction zone earthquakes near the Puerto Rico Trench can cause severe inundation at Virginia Beach. The results also highlight that large waves can arrive 12-30 hours after the first wave arrival for far-field, small to medium sized tsunamis (maximum wave amplitudes near or smaller than 1 meter). This typical long arrival time requires longer warning duration for such events. Wavelet analyses show broad and relatively long resonant periods from 0.5 hours to several hours, also suggesting the necessity for long warning durations.

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The simulated magnitude-9.3 tsunamis show an impressive local variability of tsunami amplitudes at Virginia Beach, and indicate the complexity of forecasting tsunami amplitudes at a coastal location. It is essential to use high-resolution models to provide the accuracy useful for coastal tsunami forecasts and practical guidance.

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37 **1** Introduction

38 The National Oceanic and Atmospheric Administration (NOAA) Center for Tsunami 39 Research, located at NOAA's Pacific Marine Environmental Laboratory (PMEL), has 40 developed a tsunami forecasting system for operational use by NOAA's two Tsunami 41 Warning Centers located in Hawaii and Alaska (Titov et al., 2005; Titov, 2009). The 42 forecast system combines real-time deep-ocean tsunami measurements from tsunameters 43 (González et al., 2005; Meinig et al., 2005, Bernard et al., 2006, Bernard and Titov, 2007) and the Method of Splitting Tsunami (MOST) model, a suite of finite difference 44 45 numerical codes based on the nonlinear shallow water wave equations (Titov and 46 Synolakis, 1998; Titov and González, 1997; Synolakis et al., 2008; Titov et al., 2011) to 47 produce real-time forecasts of tsunami arrival time, heights, periods and inundation. To 48 achieve accurate and detailed forecasts of tsunami impact for specific sites, high-49 resolution tsunami forecast models are under development for United States coastal 50 communities at risk (Tang et al., 2008a; 2009; 2010; Arcas and Uslu, 2010; Righi and 51 Arcas, 2010; Uslu et al. 2010; Wei and Arcas, 2010). The resolution of these models has 52 to be high enough to resolve the dynamics of a tsunami inside a particular harbor, 53 including influences of major harbor structures such as breakwaters and seawalls. These 54 models have been integrated as crucial components into the tsunami forecast system.

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56 As of March 2013 the forecast system real-time measurements come from a network 57 of 58 tsunameter stations deployed at optimal locations in the Pacific, Atlantic, Indian 58 Oceans, Caribbean Sea, the Gulf of Mexico and South China Sea (Spillane et al., 2008). 59 While the buoy array is owned and maintained by nine different nations (U.S.A., 60 Australia, Chile, China, Japan, India, Indonesia, Thailand and Russian), the data from the 61 entire array are made publically available in real-time via the Global 62 Telecommunications System (GTS). The data from the tsunameters is used to provide 63 guidance by comparing them to pre-computed open ocean model results. These pre-64 computed propagation models currently cover all three ocean basins (Pacific, Atlantic, 65 and Indian), and are comprised of 1,725 different tsunami scenarios with initial 66 deformations covering the major tsunamigenic subduction zones throughout the world 67 (Figure 1 and Table 1). High-resolution forecast inundation models are now set up for 75 68 U.S. coastal communities (e.g. Fig. 1). The fully implemented system will use real-time 69 data from the tsunameter network to provide high-resolution tsunami forecasts for at least 70 75 communities in the U.S. by 2013, with additional models envisioned later for smaller

71 communities. Since its first testing in the 17 November 2003 Rat Island tsunami, the

forecast system has produced experimental real-time forecasts for more than 20 tsunamis
 in the Pacific and Indian oceans (Titov *et al.*, 2005; Wei *et al.*, 2008; Titov, 2009; Titov

and Tang, 2011; Tang *et al.*, 2012; <u>http://nctr.pmel.noaa.gov/database_devel.html</u>). The

75 forecast method has also been tested with data from nine additional events that produced

76 deep-ocean tsunameter data including several near-field tsunamis

77 (<u>http://nctr.pmel.noaa.gov/database_devel.html</u>; Titov *et a*l., 2005; Tang *et al.*, 2008b;

78 Wei et al., 2012).

80 This report describes the development and testing of the Virginia Beach forecast 81 model. The objective in developing this model is to provide NOAA's Tsunami Warning 82 Centers the ability to assess danger posed to Virginia Beach following tsunami generation 83 in the Atlantic Ocean Basin with a goal to provide accurate and timely forecasts to enable 84 the community to respond appropriately. Dr. Aurelio Mercado developed the first version 85 of a Virginia Beach forecast model in 2007. It was updated in 2008 and had been 86 working in the tsunami forecast system. As new bathymetric/topographic and tsunami 87 data came in and the model development technique progressed further, the model had 88 been updated and re-tested in March 2013. A secondary objective of the report is to 89 explore the potential tsunami impact from earthquakes at major subduction zones in the 90 Atlantic Ocean to the city by using the developed forecast model. Wavelet analysis was 91 applied to investigate the local responses to tsunami waves.

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93 The report is organized as follows. Section 2 briefly introduces NOAA's tsunami 94 forecast method. Section 3 describes the model development. Section 4 presents the 95 results and discussion, which includes model sensitivity study to friction coefficient,

model verification, and testing for simulated tsunamis. A summary and conclusion are

- 97 provided in section 5.
- 98

99 2 Forecast Method

100 NOAA's real-time tsunami forecasting scheme is a process that comprises two steps: 101 (1) construction of a propagation scenario via inversion of deep ocean tsunameter 102 measurements with pre-computed tsunami source functions; and (2) coastal predictions 103 by running high-resolution forecast models in real time (Titov et al. 1999; 2005; Titov 104 2009; Tang et al., 2009, Tang et al., 2012). The tsunameter-constrained tsunami source, 105 the corresponding offshore scenario from the tsunami source function database, and high-106 resolution forecast models cover the entire evolution of earthquake-generated tsunamis, 107 generation, propagation and coastal inundation, providing a complete tsunami forecast 108 capability.

109

1102.1 Construction of a Propagation Scenario Based on Deep-Ocean Tsunameter111Measurements and Pre-Computed Tsunami Source Functions

112 Several real-time data sources, including seismic data, coastal tide gage and deep-113 ocean data have been used for tsunami warning and forecasting (Satake *et al.*, 2008; 114 Whitmore, 2003; Titov, 2009). NOAA's strategy for the real-time forecasting is to use 115 deep-ocean measurements at tsunameter stations, also known as DART[™] (Deep-ocean 116 Assessment and Reporting of Tsunami) buoys, as the primary data source due to several 117 key features: (1) tsunameters provide a direct measure of tsunami waves, unlike seismic 118 data, which are an indirect measure of tsunamis, (2) deep ocean tsunami measurements 119 are in general the earliest tsunami information available, since tsunamis propagate much 120 faster in deep ocean than in shallow coastal area where coastal tide gages are located, (3)

121 compared to coastal tide gages, tsunameter data with a high signal to noise ratio can be

- 122 obtained without interference from harbor and local shelf effects, and (4) wave dynamics
- of tsunami propagation in deep water is assumed to be linear (Kânoğlu and Synolakis,
 2006; Liu, 2009). This linear process allows application of efficient inversion schemes.

125 Time series of tsunami observations in deep water (depths << wave length) can be 126 decomposed into a linear combination of a set of tsunami source functions in the time 127 domain by a linear Least Squares method (Percival et al., 2011). The coefficients 128 obtained through this inversion process are called *tsunami source coefficients*. During 129 real-time tsunami forecasting, seismic waves propagate much faster than tsunami waves 130 so the initial seismic magnitude can be estimated before the tsunameter data are available. 131 Since time is of the essence, this initial tsunami forecast is based on the seismic 132 magnitude only. An updated forecast will be made via the inversion method when 133 tsunameter is available.

134

135 Titov et al. (1999; 2001) conducted sensitivity studies on far-field deep-water 136 tsunamis with different parameters of an elastic deformation model described in 137 Gusiakov (1978) and Okada (1985). The results showed source magnitude and location 138 essentially define far-field tsunami signals for a wide range of subduction zone 139 earthquakes. Other parameters have a secondary influence and can be pre-defined during 140 the forecast. Based on these results, tsunami source function databases for the Pacific, 141 Atlantic, and Indian Oceans have been built using these pre-defined source parameters: length = 100 km, width = 50 km, slip = 1 m, rake = 90 or -90 and rigidity = 4.5×10^{10} 142 143 N/m^2 . The other parameters (strike, dip, and depth) are location-specific and are chosen 144 with knowledge of the subduction zone where they are located. Details of the propagation 145 database are described in Gica et al. (2008). Each tsunami source function models a tsunami generated by a typical magnitude-7.5 earthquake with predefined source 146 147 parameters mentioned above. Figure 1 shows the locations of tsunami source functions. 148 Figure 2 shows the maximum amplitudes at Virginia Beach offshore from the tsunami 149 source functions in Atlantic Ocean.

150

151 The tsunami source functions in the database are computed with a time step of 10 152 seconds and a spatial resolution of 4-arc-minute (approximately 7.4 km along the north-153 south direction). The output (offshore wave height and depth-average velocities of the 154 entire domain) are then compressed and saved every 1 minute in time and 16-arc-minute 155 in space (Tolkova, 2007). As inundation is calculated by the high resolution forecast 156 models, the propagation scenarios do not include inundation, and a reflection boundary 157 condition is enforced at 20 m water depth (Gica et al., 2008), and friction is assumed to 158 be negligible.

159

When tsunami waves propagate into shallow water, the steady-state assumption requires no net energy losses or gains. The decrease in transport speed must be compensated by an increase in energy density in order to maintain a constant energy flux. The low spatial resolution and simplified boundary conditions of the propagation model result in inaccuracies in near-shore dynamics. As a consequence, the numerical dissipation (due to the low spatial resolution) will cause energy decay in the propagation modeling (Tang *et al.*, 2012). Based on the consideration of energy conservation, we
 have developed high-resolution, site-specific inundation forecast models built on the
 MOST model to more accurately simulate the near shore wave dynamics.

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170 That percentage of energy released from an earthquake that is transferred into the 171 water column during tsunami generation is difficult to accurately model using seismic 172 methods. However, the goal of tsunameter inversion is not to quantify the energy at the 173 initial stage of tsunami generation. Instead, we try to quantify the amount of wave energy 174 that propagates outside the source area in the form of surface long gravity waves, which 175 can be well measured by the tsunameter stations. Since it is this propagating energy that 176 results in the impacts at the coast, we estimate the tsunami source (the propagation 177 scenario) by directly measuring the deep ocean tsunami data. Regardless of the details of 178 earthquake processes for tsunami generation at the initial stage, the inversion can ensure 179 the propagation scenario gives the best approximation to the tsunami measurements, and 180 therefore, the best estimation of the total energy transferred to the tsunami waves. The 181 database can provide offshore forecasts of tsunami amplitudes and all other wave 182 parameters immediately once the inversion is complete. The tsunami source, constrained 183 by real-time tsunami measurements, provides an accurate offshore tsunami scenario 184 without additional time-consuming deep-water model runs.

185

186 **2.2** Coastal Predictions by Using High-Resolution Forecast Models in Real-Time.

187 High-resolution forecast models are designed for the final stage of the evolution of 188 tsunami waves: coastal runup and inundation. Once the tsunameter-constrained tsunami 189 source is obtained (as a linear combination of tsunami source functions), the pre-190 computed time series of offshore wave height and depth-averaged velocity from the 191 model propagation scenario are applied as the dynamic boundary conditions for the 192 forecast models. This saves the simulation time of basin-wide tsunami propagation. 193 Tsunami inundation and nearshore currents are highly nonlinear processes, therefore a 194 linear combination would not provide an accurate solution. A high-resolution model is 195 also required to resolve shorter tsunami wavelengths nearshore with accurate 196 bathymetric/topographic data. The forecast models are constructed with the Method of 197 Splitting Tsunami (MOST) model, a finite difference tsunami inundation model based on 198 the nonlinear shallow-water wave equations (Titov and Synolakis, 1998; Titov and 199 González, 1997; Synolakis et al., 2008; Titov et al., 2011). Each forecast model contains 200 three telescoping computational grids with increasing resolution, covering regional, 201 intermediate and near shore areas. Run-up and inundation are computed at the coastline. 202 The highest resolution grid includes the population center and coastal water level stations 203 for forecast verification. The grids are derived from the best available 204 bathymetric/topographic data at the time of development, and will be updated as new 205 survey data become available.

206

The forecast models are optimized for speed and accuracy. By reducing the computational areas and grid resolutions, each model is optimized to provide 4-hour event forecasting results in a maximum of 10 minutes of computational time using asingle processor, while still providing enough accuracy for forecasting. To ensure

forecast accuracy at every step of the process, the model output is validated with

historical tsunami records when available and compared to numerical results from the original full-resolution, and full-extent "reference" inundation model. In order to provide

warning guidance over a long duration during a tsunami event, each forecast model has

been developed to provide simulation output for up to 24 hours (or 30 hours for the

216 Atlantic sites) from the time of tsunami generation.

217

218 **3** Model Development

219 **3.1 Forecast area**

220 Virginia Beach is a coastal city in the mid-Atlantic region in the southeastern corner 221 of Virginia, within the geologic region called the Atlantic Coastal Plain (Tayler et al., 222 2007). The city is bordered by the Atlantic Ocean to the east and the Chesapeake Bay to 223 the north (Fig. 4). The Atlantic Coastal Plain features a thick basement layer of igneous 224 and metamorphic rock overlain with a thick wedge of sediment that increases in thickness 225 and dips towards the eastern shoreline (Tayler *et al.*, 2007). This sedimentary wedge 226 consists primarily of eroded clays, sands, and gravel from the Appalachian mountains, 227 covered with a thin layer of marine sands deposited in a series of sea level changes. 228 Chesapeake Bay also contains an impact crater estimated to be 35 million years old, 229 stretching 90 km in diameter (USGS Fact Sheet 049-98). As the plains were uplifted, 230 numerous peninsulas were incised by stream cutting, with the larger rivers forming tidal 231 rivers.

Virginia Beach's enormous popularity is derived from 28 miles of beach front which, according to the Guinness Book of Records, is the largest pleasure beach in the world. The present-day beach extent is shown in Figure 5. Since the late 1920s, the city has experienced tremendous growth both as a resort and as a center of industry for the East Coast. Virginia Beach is the largest city in Virginia. Figure 6 shows the population density for the city, with many high-density tracts right along the coveted beach front area.

The nearest Nation Ocean Service tide gage is the Chesapeake Bay Bridge Tunnel station (e.g. Allen et al., 2008), which is about 18 km northwest of the warning point for Virginia Beach. According to the station, the mean rage of tide is 0.777m. The mean high water is used as the reference level for the forecast model to provide a worst case for inundation forecast.

Although no tsunami run-up data were found for Virginia Beach from NGDC's database, due to its low lying coastal area, high coastal population density and the potential tsunami hazard from Caribbean Sea subduction zone earthquakes, Virginia Beach is in need of a forecast model to aid site-specific evacuation decisions.

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| 249 | 3.2 Bathymetry and Topography |
|--|---|
| 250 251 252 253 254 255 | In January of 2007, NOAA's National Geophysical Data Center (NGDC) developed a 1/3" digital elevation model (DEM) covering the Virginia Beach, Virginia region (Taylor et al., 2007). At the latitude of Virginia Beach, Virginia (36°51' N, 76°00' W) 1/3 arc-second of latitude is equivalent to 10.27 meters; 1/3 "of longitude is equivalent to 8.26 meters. The details of this "base" DEM development can be found in Taylor <i>et al.</i> (2007), and an overview is provided here: |
| 256 | |
| 257 | The DEM was generated from diverse digital datasets in the region (sources |
| 258 | shown in Fig. 7) and were designed to represent modern morphology. The digital data |
| 259 | were obtained from several U.S. federal, state and local agencies: |
| 260 | (1) Bathymetry data include |
| 262 | (1) Buillymenty data mende |
| 263 | • NOAA's National Ocean Service (NOS) hydrographic survey data |
| 264 | Recent NOS shallow-water multibeam survey data |
| 265 | • USACE surveys of dredged shipping channels and the Intracoastal |
| 266 | Waterway |
| 267 | NGDC-digitized Atlantic Intracoastal Waterway |
| 268 | ENC-extracted sounding data |
| 269 | |
| 270 | (2) Topography dataset include: |
| 271 | |
| 272 | City of Virginia Beach 2004 LiDAR data with ~2m spacing USCS 1000 NED 1/2 to 1 are second data |
| 273 | • USGS 1999 NED 1/3 to 1 arc-second data. |
| 274 | All dataset were shifted to the World Geodetic System 1984 (WGS84) horizontal |
| 276 | datum and transferred to the MHW vertical datum |
| 277 | |
| 278 | A grid generation algorithm was used to generate 36 arc-second and 60 arc- |
| 279 | second grids for the high-resolution reference model, covering the East coast of Virginia. |
| 280 | The data consisted largely of the above-referenced 1/3 arc-second Virginia Beach DEM, |
| 281 | but outside it's extent, data were also used from the following sources: |
| 282 | |
| 283 | Virginia Beach VA 1/3" |
| 284 | • Morehead City, NC 1/3 " |
| 285 | • Ocean City, MD 1/3 " |
| 286 | • Cape Hatteras 1/3 " |
| 287 | • Savannah GA 1/3 " |
| 288 | • Atlantic City NJ 1/3 " |
| 289 | • Nantucket MA 1/3 " |
| 290 | • Montauk NY 1/3 " |

- 291 Daytona Beach FL 1/3 " • Myrtle Beach SC 1/3 " 292 • East Coast 9 " 293 • 294 Atlantic Test 1 ' (ETOPO1 from NGDC) ٠ 295 296 The bathymetry and topography at Virginia Beach used in the development of this 297 forecast model was based on the 1/3'' DEM provided by the National Geophysical Data 298 Center. The author considers it to be a good representation of the local 299 topography/bathymetry. As new digital elevation models become available, the forecast 300 model will be updated and report updates will be posted at:
- 301
- 302 303

http://nctr.pmel.noaa.gov/forecast_reports/

304 3.3 Model Setup

305 By sub-sampling the DEMs described in section 3.2, two sets of computational 306 grids were derived for Virginia Beach, a reference inundation model and the optimized 307 forecast model.

308

The reference grids consist of four levels of telescoping grids with increasing resolution (Fig. 8). The A-grid covers the coast of Virginia in 36 arc-second. The B-grid covers Virginia Beach City, Chesapeake Bay and Pamlico Sound in 12 arc-second resolution, and run-up and inundation simulations are computed at the coastline in C grid with full 1/3 arc-second resolution.

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315 To improve the computational speed for operational use, the forecast model needs 316 to reduce the number of grid nodes, while still accurately capturing model dynamics. The 317 Virginia Beach forecast model has three levels of telescoping grids (Fig. 9). Due to the 318 shallow, wide continental shelf offshore, a 60 arc-second resolution was necessary for 319 the forecast A-grid, to propagate the wave from the propagation database (16 arc-second) 320 to the forecast site. An 18 arc-second resolution was used for the forecast B-grid, and 321 run-up and inundation simulations are computed at the coastline in C-grid at 2 arc-second 322 resolution. Figure 8c shows the Virginia Beach warning point at 284.0286°E, 6.8530°N 323 in 6.0 m of water depth.

324

325 Grid details at each level and input parameters are summarized in Table 3. As will 326 be discussed in Section 4.1, a small friction coefficient is chosen due to the shallow 327 continental shelf of U.S. East Coast. Reflection boundary conditions imposed at different 328 water depth of 0.5-10m were tested. We have noticed that even with 0.5 m, waves could 329 not propagate into some shallow areas in A- and B- grids. Due to the small friction 330 coefficient, a slightly deep offshore water depth of 5 m is chosen here to make the models 331 stable for all tested scenarios. Therefore, reflection boundary conditions were imposed at 332 5 meter water depth for the A- and B- grids

333

334 All model runs were tested on a DELL PowerEdge R510 computer equipped with 335 two Xeon E5670 processors at 2.93 Ghz, each with 12 MBytes of cache and 32GB 336 memory. The processors are hex core and support hyperthreading, resulting in the 337 computer performing as a 24 processor machine. Additionally, the testing computer 338 supports 10 Gigabit Ethernet for fast network connections. This computer configuration 339 is similar or the same as the configurations of the computers installed at the Tsunami 340 Warning Centers so the compute times should only vary slightly. For a 4-hour event 341 simulation, it takes eight processors 2 hours for the reference model while it takes only 342 \sim 8 minutes on a single processor for the forecast model.

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4 Results and Discussion

347 4.1 Sensitivity of modeled amplitude to friction coefficients

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349 Accurate simulation of tsunami induced current, run-up and inundation requires 350 high-resolution bathymetry and topography data in the run-up area and good tsunami 351 source and model parameters. Titov et al. (2005) have shown that, under these conditions, 352 the MOST run-up and inundation agree quite well with the stereo aerial photo data and 353 field survey data on Okushiri Island by the 12 July 1993 Hokkaido-Nansei-Oki Mw 7.8 354 earthquake. Wei et al. (2012) have also shown excellent agreement between the modeled 355 near-field run-up and inundation and the survey data for the March 11, 2011 Japan 356 tsunami.

357

At present, one major difficulty is the lack of high quality inundation/run-up and current measurements to verify the accuracy of topography and to calibrate the friction coefficient. Modeling work in similar near-shore geometry has suggested that lowering the friction coefficient for areas with beach slope shallower than 1:50 can have a marked affect on inundation. In this section, we focus on the sensitivity of modeled amplitude at Virginia Beach to the friction coefficient.

364

Figure 11 shows different Manning roughness coefficients (*n*) can affect the tsunami amplitude and inundations significantly at Virginia Beach for a Puerto Rico magnitude-9.3 tsunami. The testes are done using a set of testing grids with resolutions of 120, 18 and 3 arc-second for the A- B and C grids respectively. The Manning roughness coefficient (*n*) ranges from 0-0.04. The roughness coefficient can affect results not only within C-grid, but also the amplitudes in the B-grid, and those over the shallow continental shelf (depth less than 100-200m) in the A-grid.

372

The above results indicate friction does influence the results, and it is very difficult to provide the friction coefficient that is reasonable, since there are actually 375 many factors involved -- not only the roughness itself, but the approximation inherent in the shear stress of the flow, turbulent parameterization, and numerical dissipation 376 377 approximation. The Manning formula used in the MOST model is an engineering 378 empirical "roughness" estimation only. Using any specific number is not really validated 379 in any way for tsunamis and the choice of a specific coefficient for a specific site is fairly 380 arbitrary. The goal is to account for friction that we know exists, and to improve the 381 stability of the runs for a particular site. The best way to validate the friction is with 382 observation data, but such data is rarely available for east coast sites, especially for 383 inundation. For our application, we want to be conservative in choosing the coefficient. 384

So for Virginia Beach, with beach slopes sometimes approaching 1:75, we use the smallest possible friction value that produces consistent stable computations: n=0.01. It should be noted for MOST version 4, *n* can be set to different values for different grids. For example we can set n=0 for the A- and B-grids, while a relatively larger *n* can be chosen for C-grid to stabilize the model for large run-up/run-down. Model locations with steeper beach profiles and deeper shelf depths may be best run with *n* set to the more common west-coast value of n=0.025-0.03.

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393 4.2 Model verification and stability testing

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Figure 2 shows the maximum amplitude offshore at Virginia Beach from the propagation database. The Puerto Rico trench can easily be identified as the most hazardous tsunami-generating area for Virginia Beach, with large offshore amplitudes and fast arrivals. A set of nine simulated magnitude-9.3 tsunamis was selected here for further examination (Table 3). Each simulated earthquake involves 20 tsunami source functions (10 pairs) and a uniform 25-m coefficient. Both the Virginia Beach reference and forecast models were tested with the nine scenarios.

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Figure 12 show the amplitude (η) time series at the Virginia Beach warning point for the simulated magnitude-9.3 tsunamis. Table 3 summarize the η_{max} and uncertainty due to different model setup. The uncertainty are computed as:

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407
$$uncertainty = \frac{\left|\eta_{\max 2} - \eta_{\max 1}\right|}{\eta_{\max 1}} \times 100$$

408

409 where η_{max1} and η_{max2} are the maximum water surface elevation computed by the 410 reference and forecast models respectively.

411

Based on the nine scenarios, the uncertainty in η_{max} at the Virginia Beach warning point computed by the forecast model is within 22%. The largest error of 1.07 m in η_{max} among the tested tsunamis occurs in the 5th, Puerto Rico scenario, where $\eta_1 = 4.94$ m and $\eta_{\text{max2}} = 6.01$ m cm, 22% overestimated (Fig. 12.5). The arrival time of the maximum amplitudes can be 10-30 hours after the first wave. The forecast model was tested by Figure 13 shows the model is also stable for a micro tsunami (about 0.22mm in amplitude) generated by a magnitude-6.8 earthquake in the distant South Sandwich Islands subduction zone. Accurate forecasting of small tsunamis gives forecasters the confidence to issue an all-clear for small events.

424

425 Wavelet analyses were performed for the scenarios to explore peak resonant periods, 426 $T_{\rm P}$, at the Virginia Beach warning point. Figures 14 show the amplitude spectrograms. 427 The site shows relatively long and broad resonant periods from 0.5 to 4.5 hours (Fig. 15b).

428

Figure 16 shows both the reference and forecast models produced similar maximum water elevation, maximum current and inundation limit in the study area. Large maximum currents can be seen in both the reference and forecast models for many of the scenarios, especially over the shallow areas.

433

434 Tsunami waves in the study area vary significantly for the 9 magnitude-9.3 scenarios. 435 The 5th Puerto Rico scenario produces waves with amplitudes over 6 meters at the 436 Virginia Beach warning point. The inundation and run-up for these aphysical synthetic 437 test sources can be significant, but model shows stable characteristics even for these large 438 amplitude overland flows. The 4th Dominica scenario generated wave amplitudes of over 439 2 meters at Virginia Beach. These results show the complexity and high nonlinearity of 440 tsunami waves nearshore, which again demonstrate the value of high-resolution forecast 441 model for providing accurate site-specific forecast details.

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- 443

444 5 Summary and Conclusions

445 A tsunami forecast model was developed for Virginia Beach, Virginia. The 446 computational grids for the Virginia Beach forecast model were derived from the best 447 available bathymetric and topographic data sources. The forecast model is optimally 448 constructed at 2 arc-second resolution, to enable a 4 hour inundation simulation within ~8 449 minutes of computational time using a single processor. A reference inundation model of 450 higher resolution of 1/3 arc-second was also developed in parallel, to provide a modeling 451 reference for the forecast model. Both models were tested for a set of nine simulated 452 magnitude-9.3 tsunamis. One point at the Virginia Beach, 284.0286°E, 6.8530°N in 6.0 453 meters of water depth, was chosen as the warning point for the site.

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The modeled amplitude, inundation and current are sensitive to the friction coefficient. Due to the lack of data for calibration and to be on the conservative side, we use the smallest possible friction value, n=0.01, that produces consistent stable computations for both models. 465

459

466 Mega-tsunamis from the Puerto Rico subduction zone can cause appreciable 467 inundation at the site. The results also highlight the fact that, due to the broad continental 468 shelf on the east coast, the maximum wave can arrive 12-30 hours after the first wave, 469 requiring longer warning duration for such events. The simulated magnitude-9.3 tsunamis 470 show an impressive local variability of tsunami amplitudes at Virginia Beach, and 471 indicate the complexity of forecasting tsunami amplitudes at this coastal location. It is 472 essential to use high-resolution models in order to provide enough accuracy to be useful 473 for coastal tsunami forecasts and practical guidance.

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475

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Tables

| | | Source Zone | Tsunami source functions Run time | | | | |
|-----|-------|---|-----------------------------------|---------|--------|--|--|
| No. | Abbr. | Name | Line/zone | Numbers | (hour) | | |
| 1 | ACSZ | Aleutian-Alaska-Canada-Cascadia | BAZYXW | 184 | 24 | | |
| 2 | CSSZ | Central-South American | BAZYX | 382 | 30 | | |
| 3 | EPSZ | East Philippines | BA | 44 | 30 | | |
| 4 | KISZ | Kamchatka-Kuril-Japan Trench-Izu Bonin-Marianas-Yap | BAZYXW | 229 | 24 | | |
| 5 | MOSZ | Manus Ocean Convergence Boundary | BA | 34 | 24 | | |
| 6 | NVSZ | New Britain-Solomons-Vanuatu | BA | 74 | 24 | | |
| 7 | NGSZ | North New Guinea | BA | 30 | 30 | | |
| 8 | NTSZ | New Zealand-Kermadec-Tonga | BA | 81 | 24 | | |
| 9 | NZSZ | South New Zealand | BA | 14 | 30 | | |
| 10 | RNSZ | New Ryukus-Kyushu-Nankai | BA | 44 | 24 | | |
| 11 | KBSZ | Kamchatskii-Bering Source Zone | BAZ | 13 | 13 24 | | |
| | | | Subtotal: | 1129 | | | |
| 12 | ATSZ | Atlantic | BA | 214 | 36 | | |
| 13 | SSSZ | South Sandwich | BAZ | 33 | 36 | | |
| | | | Subtotal: | 247 | | | |
| 14 | IOSZ | Adaman-Nicobar-Sumatra-Java | BAZY | 307 | 24 | | |
| 15 | MKSZ | Makran | BA | 20 | 24 | | |
| 16 | WPSZ | West Philippines | BA | 22 | 24 | | |
| | | | Subtotal: | 349 | | | |
| | | | Total: | 1725 | | | |

Table 1 Tsunami source functions in the Pacific, Atlantic and Indian Oceans.

| Grid | Region | Reference | Model | | Forecast model | | |
|---------|---------------|-------------------------------------|---------------|-------------|---|------------------|-------|
| | | Coverage | Cell | Time | Coverage | Cell | Time |
| | | Lon. (°E) | Size | Step | Lon. (°E) | Size | Step |
| | | Lat. (^o N) | (") | (sec) | Lat. (^o N) | (") | (sec) |
| А | Virginia | 281.9241-289.6941 | 36 | 3.5 | 281.9224- 289.7057 | 60 | 5.2 |
| | | 32.5007-41.3507 | (778x88 | 36) | 32.6774-41.344 | (468x521) | |
| В | Virginia | 282.8707-285.6541 34.764-39.6474 | 12 (836x14 | 1.5 466) | 283.5391- 284.7141 36.1024-37.9324 283.9686- | 18 (236x367) | 5.2 |
| С | Virginia | 283.9953-284.0796 | 1/3 | 0.4 | 284.0791 | 2 | 2.6 |
| | Beach | 36.8143-36.9476 | (911x | (1441) | 59.5064-59.6647 | (200x3 | 54) |
| Minim | um offshore | denth (m) | 5 | | | 5 | |
| Water | denth for dr | uland (m) | 0.1 | | 0.1 | | |
| Triatia | | (m^2) | 0.1 | 0.0001 | 0.1 | | |
| rncu0 | ii coemcient | . (11.) | $\sim 2 hou$ | u.0001 | t | 0.0001 | |
| Compu | itational tim | e for a 4-hr simulation | 8 proc | cessors | 7.86 min u | sing 1 processor | r |

 Table 2 MOST setups for the Virginia Beach reference and forecast models.

Table 3 Sources of the 9 simulated magnitude-9.3 tsunamis and the maximum computed wave crests at the Virginia Beach warning point.

| No | . Subo Zone | d. \$ ∋ | Souro | ce i | alpha | Ref eta (I | E. mo amax – n) (1 | del tmax hour) | Fore etam (m) | ecast Mo nax tmax (hour) | del H : (T | Error n) (%) | Location |
|----|----------------|------------|-------|------|-------|---------------------|--------------------------|----------------------|-----------------------|--------------------------------|----------------------|-----------------|--------------------|
| | | | | | | | | | | | | | |
| 1 | atsz | AB | 1- | 10 | 25 | 0.45 | 15.7 | 74 (| 0.45 | 31.487 | -0.00 | 0 0 | Panama |
| 2 | atsz | AB | 12- | 21 | 25 | 0.53 | 16.6 | 41 (| 0.49 | 16.680 | -0.05 | 5 -9 | Colombia |
| 3 | atsz | AB | 22- | 31 | 25 | 0.96 | 17.8 | 01 (| 0.86 | 17.805 | -0.09 | 9 -10 | Venezuela |
| 4 | atsz | AB | 38- | 47 | 25 | 2.52 | 4.7 | 53 2 | 2.47 | 5.277 | -0.05 | 5 -2 | Dominica |
| 5 | atsz | AB | 48- | 57 | 25 | 4.94 | 4.4 | 73 (| 6.01 | 4.429 | 1.07 | 7 22 | Puerto Rico |
| 6 | atsz | AB | 58- | 67 | 25 | 1.06 | 4.6 | 63 | 1.09 | 4.704 | 0.03 | 3 3 | Cayman |
| 7 | atsz | AB | 68- | 77 | 25 | 0.23 | 26.0 | 85 (| 0.23 | 26.002 | 0.01 | L 0 | Gulf of Honduras |
| 8 | atsz | AB | 82- | 91 | 25 | 1.33 | 5.0 | 60 | 1.38 | 5.095 | 0.04 | 1 3 | U.S. Virgin Is. |
| 9 | SSSZ | AB | 1- | 10 | 25 | 0.63 | 28.4 | 60 (| 0.58 | 28.471 | -0.05 | 5 -7 | South Sandwich Is. |

Appendix A.

Since the initial development of the forecast model for Virginia Beach, Virginia, the parameters for the input file for running the forecast and reference models have been changed to reflect changes to the MOST model code. The following appendix lists the new input files for Virginia Beach.

A1. Reference model *.in file for Virginia Beach, Virginia—updated for 2013

```
# ------ MOST Run 1 ------
# 0. Preparations
echo '#-----#'
echo '# Preprocess MOST input #'
echo '#-----#'
set main_dir="/home/tg23/data/tang/sims/virginiabeach/"
set np="8"
setenv OMP_NUM_THREADS $np
set path_w="$main_dir/virgv4_S03_at_ab22T31rb2_c1_05m_fp01_19h/"
set path e="most4"
set path_src="/grid/tg23/data/tang/src_nc/src_sim_test/virg/S03_at_ab22T31_virg_"
if ( -d $path_w ) then
echo $path_w 'exist'
 cd $path_w
else
 echo Creating directory $path_w
 mkdir $path_w
 cd $path_w
endif
ln -sf /home/tg23/data/tang/bathy/virginiabeach/virg_rb2//*.nc .
# -----
# 1. Generate INPUT for MOST
cat > most3 facts nc.inA<< EOF
0.005 Minimum amplitude of input offshore wave (m):
5
    Input minimum depth for offshore (m)
0.1 Input "dry land" depth for inundation (m)
0.0001 Input friction coefficient (n^{**}2)
2
    Number of grids
2
     Interpolation domain for outer boundary
2
     inner boundary
RA_VirginiaBeach_36ss_20130211.nc
RB_VirginiaBeach_12s_20130211.nc
1
      Runup flag
```

Input time step (sec) 3.5 19543 Input amount of steps 0 **COntunue** after input stops 9 Input number of steps between snapshots 1 saving inner boundaries every n-th timestep ...Saving grid every n-th node, n= 1 1=initial deformation 0 EOF cp most3_facts_nc.inA most3_facts_nc.in #\$path_e A \$path_src most3_facts_nc.in cat > most3 facts nc.inB<< EOF 0.005 Minimum amplitude of input offshore wave (m): Input minimum depth for offshore (m) 5 Input "dry land" depth for inundation (m) 0.1 0.0001 Input friction coefficient (n**2) 2 Number of grids 2 Interpolation domain for outer boundary 2 inner boundary RB_VirginiaBeach_12s_20130211.nc RC_Virginia_Beach_1_3s_c2_NGDC.nc Runup flag 1 1.5 Input time step (sec) 45600 Input amount of steps COntunue after input stops 0 20 Input number of steps between snapshots 1 saving inner boundaries every n-th timestep 1 ...Saving grid every n-th node, n= 0 1=initial deformation EOF cp most3_facts_nc.inB most3_facts_nc.in \$path_e B A most3_facts_nc.in cat > most3_facts_nc.inC<< EOF 0.005 Minimum amplitude of input offshore wave (m): -300 Input minimum depth for offshore (m) 0.1 Input "dry land" depth for inundation (m) 0.0001 Input friction coefficient (n**2) 1 Number of grids 2 Interpolation domain for outer boundary 2 inner boundary RC_Virginia_Beach_1_3s_c2_NGDC.nc 2 Runup flag 0.38 Input time step (sec) 180000 Input amount of steps 0 COntunue after input stops 79 Input number of steps between snapshots saving inner boundaries every n-th timestep 1 1 ...Saving grid every n-th node, n= 0 1=initial deformation EOF cp most3_facts_nc.inC most3_facts_nc.in \$path_e C B most3_facts_nc.in

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A2. Forecast model *.in file for Virginia Beach, Virginia—updated for 2013

| 0.0001 | Minimum amplitude of input offshore wave (m) | | | | | |
|--|--|---------------------|--|--|--|--|
| 5 | Input minimum depth for offshore (m) | | | | | |
| 0.1 Input "dry land" depth for inundation (m | | | | | | |
| 0.0001 | Input friction coefficient (n* | **2) | | | | |
| 1 | runup flag for grids A and B | (1=yes,0=no) | | | | |
| 300.0 blowup | limit | | | | | |
| 2.6 | Input time step (sec) | | | | | |
| 16615 | Input amount of steps | | | | | |
| 2 | Compute "A" arrays every n- | -th time step, n= | | | | |
| 2 | Compute "B" arrays every n- | -th time step, n= | | | | |
| 10 | Input number of steps betwee | een snapshots | | | | |
| 0 | Starting from | | | | | |
| 1 | Saving grid every n-th node | e, n= | | | | |
| FA_Virg_Beach_6 | 50s_20130211.ssl | | | | | |
| FB_Virg_Beach_2 | 18s_20130211.ssl | | | | | |
| FC_Virg_Beach_2 | 2s_c7_NGDC.ssl | | | | | |
| /grid/tg23/data | /tang/src_nc/src_sim_test/v | irg/ | | | | |
| ./ | | | | | | |
| 1111 | NetCDF output for A, B, C, SI | FT | | | | |
| 1 | Timeseries locations: | | | | | |
| 3 109 223 | Virginia Beach 284.0286E | 36.8530N, 6 m depth | | | | |

Figures

| Figure 14 (a) Modeled η time series at Virginia Beach warning point for the nine simulated |
|--|
| magnitude-9.3 tsunamis. (b) Wavelet–derived amplitude spectrogram for the reference model. (c and d) Real part of the spectrograms computed by the reference and forecast models |
| Figure 15 (a) Forecast uncertainty in the η_{max} at the Virginia Beach warning point. (b) Uncertainty v.s. peak period. η_{max1} and T_{p1} , maximum water elevation and peak period at the warning point from the reference model. η_{max2} and T_{p2} , maximum water surface elevation and peak period at the |



Figure 1 (a) Overview of the tsunami Forecast System. System components include tsunameter (DART) network (yellow triangles), pre-computed tsunami source function (unfilled black rectangles) and high-resolution forecast models (red squares). Filled color shows the computed offshore maximum sea surface elevation in m for a simulated magnitude-9.3 Caribbean tsunami (#5). Contours indicate the travel time in hours.



Figure 2 Maximum sea surface elevation (η_{max}) offshore Virginia Beach from 214 magnitude-7.5 earthquakes. Data were taken from NCTR's pre-computed propagation database for Atlantic Ocean. Heights and colors of the bars represent η_{max} and the first arrival at the offshore point (74.1333°W, 36.9461°N; water depth = 2269 m) respectively. 1-9, locations for nine simulated magnitude-9.3 tsunamis.



Figure 3 Historical tsunamis in Atlantic Ocean and Caribbean Sea (National Geographic Data Center's database).





Figure 4 NOAA charts, (a) 13003 and (b) 12208, show Virginia Beach Soundings in fathoms at Mean Lower Low Water. Contour and summit elevation values are in feet above Mean Sea Level.



Imagery @2013 Commonwealth of Virginia, DigitalGlobe, GeoEye, U.S. Geological Survey, USDA Farm Service Agency -



Figure 5 Aerial photos of Virginia Beach.


Figure 6 Population density at Virginia Beach (2000 Census).



Figure 7 Bathymetric and topographic data source overview for the high-resolution Virginia Beach DEM. Image courtesy of Tayler *et al.* (2007).







Figure 8 Grid setup for the Virginia Beach reference model. Resolutions are (a) 36", (b) 12" and(c) 1/3". Red boxes are boundaries of the telescoped grids for the reference model.









Figure 9 Grid setup for the Virginia Beach forecast model. Grid resolutions are (a) 60 ", (b) 18", and (c) 2". Red boxes, boundaries of the telescoping grids. Red dot, Virginia Beach warning point (284.0286°E, 36.8530°N; water depth= 6m).



Figure 10 Sensitivity of η to friction coefficients. Results were computed by a set of testing grids (Fig. 11) for a magnitude 9.3 Caribbean tsunami.



Figure 11 Sensitivity of η_{max} to friction coefficients. Results were computed by a set of testing grids for a magnitude-9.3 Caribbean tsunami.





Figure 12 Modeled η time series by the Virginia Beach reference and forecast models for simulated magnitude-9.3 tsunamis.



Figure 13 Modeled η time series computed by the Virginia Beach forecast model for a simulated micro tsunami. The tsunami was generated from a Magnitude 6.8 earthquake from South Sandwich Islands Subduction (0.1 × B11).











Figure 14 (a) Modeled η time series at Virginia Beach warning point for the nine simulated magnitude-9.3 tsunamis. (b) Wavelet–derived amplitude spectrogram for the reference model. (c and d) Real part of the spectrograms computed by the reference and forecast models.



Figure 15 (a) Forecast uncertainty in the η_{max} at the Virginia Beach warning point. (b) Uncertainty v.s. peak period. η_{max1} and T_{p1} , maximum water elevation and peak period at the warning point from the reference model. η_{max2} and T_{p2} , maximum water surface elevation and peak period at the warning point computed by the forecast model.







(1) ATSZ AB1-10R Mw 9.3



























(5) ATSZ AB48-57 Mw 9.3















Figure 16 Maximum water elevation and current computed by the Virginia Beach reference and forecast models for the simulated magnitude-9.3 tsunamis.

Appendix B

Propagation Database: Atlantic Ocean Unit Sources





| Segme | ent | Description | Lon | $gitude(^{o}E)$ | $Latitude(^{o}N)$ | $Strike(^{o})$ | $\operatorname{Dip}(^{\mathrm{o}})$ | Depth (km) |
|----------------------|----------------------|---------------|--------------|-----------------|-------------------|----------------|-------------------------------------|--------------|
| atsz–1a | Atla | ntic Source / | Zone | -83.2020 | 9.1449 | 120 | 27. | 5 28.09 |
| atsz–1b | Atla | ntic Source ' | Zone | -83.0000 | 9.4899 | 120 | 27. | 5 5 |
| atsz–2a | Atla | ntic Source ' | Zone | -82,1932 | 8.7408 | 105.1 | 27. | 5 28.09 |
| atsz–2b | Atla | ntic Source ' | Zone | -82.0880 | 9.1254 | 105.1 | 27. | 5 5 |
| atsz-3a | Atla | ntic Source ' | Zone | -80.9172 | 9.0103 | 51.31 | 30 | 30 |
| atsz–3b | Atla | ntic Source ' | Zone | -81.1636 | 9.3139 | 51.31 | 30 | 5 |
| atsz-4a | Atla | ntic Source ' | Zone | -80.3265 | 9.4308 | 63.49 |) 30 | 30 |
| atsz–4b | Atla | ntic Source ' | Zone | -80.5027 | 9.7789 | 63.49 |) 30 | 5 |
| atsz–5a | Atla | ntic Source ' | Zone | -79.6247 | 9.6961 | 74.44 | 1 30 | 30 |
| atsz–5b | Atla | ntic Source ' | Zone | -79.7307 | 10.0708 | 74.44 | 1 30 | 5 |
| atsz–6a | Atla | ntic Source ' | Zone | -78.8069 | 9.8083 | 79.71 | 30 | 30 |
| atsz–6b | Atla | ntic Source ' | Zone | -78.8775 | 10,1910 | 79.71 | 30 | 5 |
| atsz-7a | Atla | ntic Source ' | Zone | -78.6237 | 9,7963 | 127.2 | 2 30 | 30 |
| atsz–7b | Atla | ntic Source ' | Zone | -78.3845 | 10,1059 | 127.2 | 2 30 | 5 |
| atsz-8a | Atla | ntic Source ' | Zone | -78.1693 | 9.3544 | 143.8 | 3 30 | 30 |
| atsz–8b | Atla | ntic Source ' | Zone | -77.8511 | 9.5844 | 143.8 | 3 30 | 5 |
| atsz-9a | Atla | ntic Source ! | Zone | -77 5913 | 8 5989 | 139.0 |) 30 | 30 |
| atsz–9h | Atla | ntic Source ' | Zone | -77 2900 | 8 8493 | 139.0 |) 30 | 5 |
| atsz-10a | Atla | ntic Source ' | Zone | -75 8109 | 9.0881 | 4 67 | , 50 17 | 19 62 |
| atsz–10b | Atla | ntic Source ' | Zone | -76 2445 | 9 1231 | 4 67 | 17 | 5 |
| atsz-11a | Atla | ntic Source ' | Zone | -75 7406 | 9 6929 | 19.65 | 7 17 | 19 62 |
| atsz 11a | Atla | ntic Source ' | Zone | -76 1511 | 9.0929 | 19.07 | 7 17 | 15.02 |
| atsz = 12a | Atla | ntic Source ' | Zone | -75 4763 | 10 2042 | 40.4 | 17 | 19 62 |
| atsz 12a atsz–19b | Atla | ntic Source ' | Zone | -75 8080 | 10.2042 | 40.4 | 17 | 15.02 |
| atsz 120 | Atla | ntic Source / | Zono | 74 0014 | 10.4020 | 40.4 | 7 17 | 10.62 |
| atsz-13a | Atla | ntic Source / | Zono | -74.3314 | 11 1064 | 47.17 | 7 17 | 19.02 |
| atsz=130 | Atla | ntic Source ' | Zone | -74 5666 | 11.1004 | 71.68 | 17 | 10.62 |
| atsz 14a | Atla | ntic Source ' | Zone | -74.7043 | 11.0708 | 71.00 | 17 | 15.02 |
| $a_{1}s_{2}-140$ | Atla | ntic Source ' | Zone | -73 4576 | 11.4780 | 11.00 | $\frac{17}{17}$ | 10.62 |
| atsz 15a | Atla | ntic Source / | Zono | 73 7805 | 12 0024 | 42.03 | $\frac{17}{17}$ | 15.02 |
| atsz-150 | Atla | ntic Source / | Zono | -13.1803 | 12.0324 | 42.03 54.75 | ז ז ג ג 17 | 10.62 |
| atsz-10a | Atla | ntic Source / | Zono | 73 2320 | 12.5505 | 54.70 |) 17 (17 | 19.02 |
| atsz-100 | Atla | ntic Source / | Zono | -73.2323 | 12.0875 | 94.70 81.06 | 3 17 | 10.62 |
| atsz-17a | Atla | ntic Source / | Zone | -72.0404 | 12.0001 | 81.90 |) 17 3 17 | 19.02 |
| atsz-170 | Atla | ntic Source . | Zone | -72.0071 | 12.9314 | 70.65 | 17 | 10.62 |
| atsz-18h | Atla | ntic Source / | Zono | -71.6830 | 12.0174 | 79.00 | 17 | 19.02 |
| atsz-100 | Atla | ntic Source / | Zono | -71.0839 | 19 7078 | 86.30 | 17 17 | 10.62 |
| atsz-19a | Atla | | Zone | -10.1910 | 12.1010 | 00.34 96.35 | 5 17) 17 | 19.02 |
| atsz-190 | Atla | ntic Source . | Zone | -70.8233 | 10.1004 | 00.32 | 5 17 1 17 | 10.62 |
| atsz-20a | Atla | ntic Source . | Zone | -70.0240 | 12.7165 | 95.94 | E 17 1 17 | 19.02 |
| atsz-200 | Atla | ntic Source . | Zone | -09.9769 | 10.1407 | 95.94 | E 17 1 17 | 10.62 |
| atsz-21a | Atla | | Zone | -09.1244 | 12.0520 | 95.94 | E 17 | 19.02 |
| atsz-210 | Atla | ntic Source . | Zone | -09.0700 | 15.0592 | 90.94 | | 0 17.04 |
| atsz-22a | Atla | ntic Source . | Zone | -00.0330 | 11.4200 | 200.8 |) 10) 15 | 17.94 |
| atsz-220 | Atla | ntic Source | Zone | -08.0102 | 10.9954 | 200.8 | 9 15) 17 | 0 17.04 |
| atsz-23a | Atla | ntic Source | Zone | -07.1240 | 11.4487 | 200.8 | 9 15 \ 15 | 17.94 |
| atsz-23D | Atla | ntic Source | Zone | -07.1010 | 11.0100 | 200.8 | | 0 17.04 |
| atsz-24a | Atla | ntic Source | Zone | -66.1656 | 11.5055 | 273.3 | 5 15) 15 | 17.94 |
| atsz-240 | Atla | nuc Source | Zone Zone | -00.1911 | 11.0724 | 213.3 |) 15 1 17 | 5 17.04 |
| atsz-20a | Atla | nuc Source | Zone | -05.2120 | 11.4240 | 270.4 | 10 I | 17.94 |
| atsz-25b | Atla | nuic Source | ∠one Z | -05.2010 | 10.9934 | 276.4 | 15 15 | 6 17.04 |
| atsz-26a | Atla | ntic Source | Zone Zone | -64.3641 | 11.3516 | 272.9 | 9 15) 17 | 17.94 |
| atsz-26b | Atla | ntic Source | Zone Zone | -64.3862 | 10.9183 | 272.9 | 9 15) 17 | 5 |
| atsz-27a | Atla | ntic Source 2 | Zone | -63.4472 | 11.3516 | 272.9 | <i>i</i> 15 | 17.94 |
| | | | | | | C | ontinued of | on next page |

Table B.1: Earthquake parameters for Atlantic Source Zone unit sources.

| S | egment | Description | $Longitude(^{o}E)$ | $Latitude(^{o}N)$ | $\operatorname{Strike}(^{\mathrm{o}})$ | Dip(°) | Depth (km) |
|--------|--------|-----------------|--------------------|-------------------|--|--------|------------|
| atsz-2 | 7b At | lantic Source Z | one -63.4698 | 10.9183 | 272.9 |) 15 | 5 |
| atsz-2 | 8a At | lantic Source Z | one -62.6104 | 11.2831 | 271. | 15 | 17.94 |
| atsz-2 | 8b At | lantic Source Z | one -62.6189 | 10.8493 | 271.1 | l 15 | 5 |
| atsz-2 | 9a At | lantic Source Z | one -61.6826 | 11.2518 | 271.6 | 3 15 | 17.94 |
| atsz-2 | 9b At | lantic Source Z | one -61.6947 | 10.8181 | 271.6 | 5 15 | 5 |
| atsz-3 | 0a At | lantic Source Z | one -61.1569 | 10.8303 | 269 | 15 | 17.94 |
| atsz–3 | ob At | lantic Source Z | one -61.1493 | 10.3965 | 269 | 15 | 5 |
| atsz–3 | la At | lantic Source Z | one -60.2529 | 10.7739 | 269 | 15 | 17.94 |
| atsz–3 | 1b At | lantic Source Z | one -60.2453 | 10.3401 | 269 | 15 | 5 |
| atsz–3 | 2a At | lantic Source Z | one -59.3510 | 10.8123 | 269 | 15 | 17.94 |
| atsz–3 | 2b At | lantic Source Z | one -59.3734 | 10.3785 | 269 | 15 | 5 |
| atsz–3 | 3a At | lantic Source Z | one -58.7592 | 10.8785 | 248.6 | 6 15 | 17.94 |
| atsz–3 | 3b At | lantic Source Z | one -58.5984 | 10.4745 | 248.6 | 6 15 | 5 |
| atsz–3 | 4a At | lantic Source Z | one -58.5699 | 11.0330 | 217.2 | 2 15 | 17.94 |
| atsz–3 | 4b At | lantic Source Z | one -58.2179 | 10.7710 | 217.2 | 2 15 | 5 |
| atsz–3 | 5a At | lantic Source Z | one -58.3549 | 11.5300 | 193.7 | 7 15 | 17.94 |
| atsz–3 | 5b At | lantic Source Z | one -57.9248 | 11.4274 | 193.7 | 7 15 | 5 |
| atsz–3 | 6a At | lantic Source Z | one -58.3432 | 12.1858 | 177.7 | 7 15 | 17.94 |
| atsz–3 | 6b At | lantic Source Z | one -57.8997 | 12.2036 | 177.7 | 7 15 | 5 |
| atsz–3 | 7a At | lantic Source Z | one -58.4490 | 12.9725 | 170.7 | 7 15 | 17.94 |
| atsz–3 | 7b At | lantic Source Z | one -58.0095 | 13.0424 | 170.7 | 7 15 | 5 |
| atsz–3 | 8a At | lantic Source Z | one -58.6079 | 13.8503 | 170.2 | 2 15 | 17.94 |
| atsz–3 | 8b At | lantic Source Z | one -58.1674 | 13.9240 | 170.2 | 2 15 | 5 |
| atsz–3 | 9a At | lantic Source Z | one -58.6667 | 14.3915 | 146.8 | 3 15 | 17.94 |
| atsz–3 | 9b At | lantic Source Z | one -58.2913 | 14.6287 | 146.8 | 3 15 | 5 |
| atsz-3 | 9v At | lantic Source Z | one -59.4168 | 13.9171 | 146.8 | 3 15 | 43.82 |
| atsz-3 | 9z At | lantic Source Z | one -59.0415 | 14.1543 | 146.8 | 3 15 | 30.88 |
| atsz-4 | 0a At | lantic Source Z | one -59.1899 | 15.2143 | 156.2 | 2 15 | 17.94 |
| atsz-4 | 0b At | lantic Source Z | one -58.7781 | 15.3892 | 156.2 | 2 15 | 5 |
| atsz-4 | 0v At | lantic Source Z | one -60.0131 | 14.8646 | 156.2 | 2 15 | 43.82 |
| atsz-4 | 0z At | lantic Source Z | one -59.6012 | 15.0395 | 156.2 | 2 15 | 30.88 |
| atsz-4 | la At | lantic Source Z | one -59.4723 | 15.7987 | 146.3 | 3 15 | 17.94 |
| atsz-4 | 1b At | lantic Source Z | one -59.0966 | 16.0392 | 146.3 | 3 15 | 5 |
| atsz-4 | lv At | lantic Source Z | one -60.2229 | 15.3177 | 146.3 | 3 15 | 43.82 |
| atsz-4 | 1z At | lantic Source Z | one -59.8473 | 15.5582 | 146.3 | 3 15 | 30.88 |
| atsz-4 | 2a At | lantic Source Z | one -59.9029 | 16.4535 | 137 | 15 | 17.94 |
| atsz-4 | 2b At | lantic Source Z | one -59.5716 | 16.7494 | 137 | 15 | 5 |
| atsz-4 | 2y At | lantic Source Z | one -60.5645 | 15.8616 | 137 | 15 | 43.82 |
| atsz-4 | 2z At | lantic Source Z | one -60.2334 | 16.1575 | 137 | 15 | 30.88 |
| atsz-4 | 3a At | lantic Source Z | one -60.5996 | 17.0903 | 138.7 | 7 15 | 17.94 |
| atsz-4 | 3b At | lantic Source Z | one -60.2580 | 17.3766 | 138.7 | 7 15 | 5 |
| atsz-4 | 3y At | lantic Source Z | one -61.2818 | 16.5177 | 138.7 | 7 15 | 43.82 |
| atsz-4 | 3z At | lantic Source Z | one -60.9404 | 16.8040 | 138.7 | 7 15 | 30.88 |
| atsz-4 | 4a At | lantic Source Z | one -61.1559 | 17.8560 | 141.1 | l 15 | 17.94 |
| atsz-4 | 4b At | lantic Source Z | one -60.8008 | 18.1286 | 141.1 | l 15 | 5 |
| atsz-4 | 4y At | lantic Source Z | one -61.8651 | 17.3108 | 141.1 | l 15 | 43.82 |
| atsz-4 | 4z At | lantic Source Z | one -61.5102 | 17.5834 | 141.1 | l 15 | 30.88 |
| atsz-4 | 5a At | lantic Source Z | one -61.5491 | 18.0566 | 112.8 | 3 15 | 17.94 |
| atsz-4 | 5b At | lantic Source Z | one -61.3716 | 18.4564 | 112.8 | 3 15 | 5 |
| atsz-4 | 5y At | lantic Source Z | one -61.9037 | 17.2569 | 112.8 | 3 15 | 43.82 |
| atsz-4 | 5z At | lantic Source Z | one -61.7260 | 17.6567 | 112.8 | 3 15 | 30.88 |
| atsz-4 | 6a At | lantic Source Z | one -62.4217 | 18.4149 | 117.9 |) 15 | 17.94 |
| atsz-4 | 6b At | lantic Source Z | one -62.2075 | 18.7985 | 117.9 |) 15 | 5 |
| atsz-4 | 6y At | lantic Source Z | one -62.8493 | 17.6477 | 117.9 |) 15 | 43.82 |
| atsz-4 | 6z At | lantic Source Z | one -62.6352 | 18.0313 | 117.9 |) 15 | 30.88 |

Table B.1 – continued from previous page

Continued on next page

| Segme | nt Description | Longitude(°E) | $Latitude(^{o}N)$ | $\operatorname{Strike}(^{\mathrm{o}})$ | Dip(°) | Depth (km) |
|----------|-----------------|---|------------------------|--|-----------|------------|
| atsz–47a | Atlantic Source | Zone -63.164 | 9 18.7844 | 110.5 | 20 | 22.1 |
| atsz-47b | Atlantic Source | Zone -63.008 | 7 19.1798 | 110.5 | 20 | 5 |
| atsz–47y | Atlantic Source | Zone -63.477 |) 17.9936 | 110.5 | 20 | 56.3 |
| atsz-47z | Atlantic Source | Zone -63.320 | 5 18.3890 | 110.5 | 20 | 39.2 |
| atsz-48a | Atlantic Source | Zone -63.880 |) 18.8870 | 95.37 | 20 | 22.1 |
| atsz-48b | Atlantic Source | Zone -63.838 | 2 19.3072 | 95.37 | 20 | 5 |
| atsz-48y | Atlantic Source | Zone -63.964 | 3 18.0465 | 95.37 | 20 | 56.3 |
| atsz-48z | Atlantic Source | Zone -63.921 | 6 18.4667 | 95.37 | 20 | 39.2 |
| atsz-49a | Atlantic Source | Zone -64.815 | 3 18.9650 | 94.34 | 20 | 22.1 |
| atsz-49b | Atlantic Source | Zone -64.781 | 19.3859 | 94.34 | 20 | 5 |
| atsz-49y | Atlantic Source | Zone -64.884 |) 18.1233 | 94.34 | 20 | 56.3 |
| atsz-49z | Atlantic Source | Zone -64.849 | 2 18.5442 | 94.34 | 20 | 39.2 |
| atsz-50a | Atlantic Source | Zone -65.692 | 1 18.9848 | 89.59 | 20 | 22.1 |
| atsz-50b | Atlantic Source | Zone -65.695 | 3 19.4069 | 89.59 | 20 | 5 |
| atsz-50y | Atlantic Source | Zone -65.687 | 1 18.1407 | 89.59 | 20 | 56.3 |
| atsz-50z | Atlantic Source | Zone -65.688 | 7 18.5628 | 89.59 | 20 | 39.2 |
| atsz-51a | Atlantic Source | Zone -66.574 | 2 18.9484 | 84.98 | 20 | 22.1 |
| atsz-51b | Atlantic Source | Zone -66.613 | 3 19.3688 | 84.98 | 20 | 5 |
| atsz-51y | Atlantic Source | Zone -66.497 | 7 18.1076 | 84.98 | 20 | 56.3 |
| atsz-51z | Atlantic Source | Zone -66.535 | 3 18.5280 | 84.98 | 20 | 39.2 |
| atsz-52a | Atlantic Source | Zone -67.541 | 2 18.8738 | 85.87 | 20 | 22.1 |
| atsz-52b | Atlantic Source | Zone -67.573 | 1 19.2948 | 85.87 | 20 | 5 |
| atsz-52y | Atlantic Source | Zone -67.478 | 1 18.0319 | 85.87 | 20 | 56.3 |
| atsz-52z | Atlantic Source | Zone -67.509 | 18.4529 | 85.87 | 20 | 39.2 |
| atsz-53a | Atlantic Source | Zone -68.454 | 7 18.7853 | 83.64 | 20 | 22.1 |
| atsz-53b | Atlantic Source | Zone -68.504 | 2 19.2048 | 83.64 | 20 | 5 |
| atsz-53y | Atlantic Source | Zone -68.357 | 5 17.9463 | 83.64 | 20 | 56.3 |
| atsz-53z | Atlantic Source | Zone -68.405 | 5 18.3658 | 83.64 | 20 | 39.2 |
| atsz–54a | Atlantic Source | Zone -69.674 |) 18.8841 | 101.5 | 20 | 22.1 |
| atsz–54b | Atlantic Source | Zone -69.584 | 5 19.2976 | 101.5 | 20 | 5 |
| atsz–55a | Atlantic Source | Zone -70.704 | 5 19.1376 | 108.2 | 20 | 22.1 |
| atsz–55b | Atlantic Source | Zone -70.564 | 7 19.5386 | 108.2 | 20 | 5 |
| atsz–56a | Atlantic Source | Zone -71.536 | 3 19.3853 | 102.6 | 20 | 22.1 |
| atsz–56b | Atlantic Source | Zone -71.438 | 5 19.7971 | 102.6 | 20 | 5 |
| atsz–57a | Atlantic Source | Zone -72.353 | 5 19.4838 | 94.2 | 20 | 22.1 |
| atsz–57b | Atlantic Source | Zone -72.320 | 5 19.9047 | 94.2 | 20 | 5 |
| atsz–58a | Atlantic Source | Zone -73.158 |) 19.4498 | 84.34 | 20 | 22.1 |
| atsz–58b | Atlantic Source | Zone -73.202 | 2 19.8698 | 84.34 | 20 | 5 |
| atsz–59a | Atlantic Source | Zone -74.356 | 20.9620 | 259.7 | 20 | 22.1 |
| atsz–59b | Atlantic Source | Zone -74.276 | 1 20.5467 | 259.7 | 20 | 5 |
| atsz–60a | Atlantic Source | Zone -75.238 | 20.8622 | 264.2 | 15 | 17.94 |
| atsz–60b | Atlantic Source | Zone -75.191 | 20.4306 | 264.2 | 15 | 5 |
| atsz–61a | Atlantic Source | Zone -76.238 | 3 20.7425 | 260.7 | 15 | 17.94 |
| atsz-61b | Atlantic Source | Zone -76.163 | 20.3144 | 260.7 | 15 | 5 17.04 |
| atsz–62a | Atlantic Source | Zone -//.202 | 1 20.5910 | 259.9 | 10 | 17.94 |
| atsz-62b | Atlantic Source | Zone -77.121 | 4 20.1638 | 259.9 | 15 | 5 17.04 |
| atsz–63a | Atlantic Source | Zone -/8.154 | 20.4189 | 209 | 10 | 17.94 |
| atsz-03D | Atlantic Source | Zone -78.000 Zone 70.005 | 1 	 19.9930 | 209 | 10 | 0 17.04 |
| atoz 641 | Atlantic Source | $Z_{0110} = -79.095$ | 20.2498 | 209.2 | 10 | 11.94 |
| atsz-04D | Atlantic Source | Zone -19.009 | 5 19.8230 5 50.0779 | 209.2 | 15 | 0 17.04 |
| atsz-00a | Atlantic Source | $Z_{0110} = -60.039$ | 20.0773 10.6716 | 208.9 050 0 | 10 | 17.94 E |
| atsz-00D | Atlantic Source | $Z_{0110} = -79.950$ | 2 19.0010 5 10.0002 | 208.9 059 G | 15 | 0 17.04 |
| atsz-00a | Atlantic Source | Z_{0} R_{0} R_{0} R_{0} R_{0} | 5 19.0993 S 10.4740 | 200.0 050 C | 10 1 E | 11.94 E |
| atsz-00D | Atlantic Source | 20110 - 30.370 | 5 19.4740 5 10.7014 | 208.0 059 5 | 15 | 0 17.04 |
| atsz-0/a | Anamuc Source | Zone -81.900 | 19.7214 | 208.5 | 15 | 17.94 |

Table B.1 – continued from previous page

Continued on next page

| Segment | Description | $Longitude(^{o}E)$ | $Latitude(^{o}N)$ | $\operatorname{Strike}(^{\mathrm{o}})$ | Dip(°) | Depth (km) |
|--------------|----------------|---|--------------------|--|--------|------------|
| atsz–67b A | tlantic Source | Zone -81.8149 | 19.2962 | 258.5 | 15 | 5 |
| atsz–68a A | tlantic Source | Zone -87.8003 | 15.2509 | 62.69 | 15 | 17.94 |
| atsz–68b A | tlantic Source | Zone -88.0070 | 15.6364 | 62.69 | 15 | 5 |
| atsz–69a A | tlantic Source | Zone -87.0824 | 15.5331 | 72.73 | 15 | 17.94 |
| atsz–69b A | tlantic Source | Zone -87.2163 | 15.9474 | 72.73 | 15 | 5 |
| atsz–70a A | tlantic Source | Zone -86.1622 | 15.8274 | 70.64 | 15 | 17.94 |
| atsz–70b A | tlantic Source | Zone -86.3120 | 16.2367 | 70.64 | 15 | 5 |
| atsz–71a A | tlantic Source | Zone -85.3117 | 16.1052 | 73.7 | 15 | 17.94 |
| atsz–71b A | tlantic Source | Zone -85.4387 | 16.5216 | 73.7 | 15 | 5 |
| atsz–72a A | tlantic Source | Zone -84.3470 | 16.3820 | 69.66 | 15 | 17.94 |
| atsz–72b A | tlantic Source | Zone -84.5045 | 16.7888 | 69.66 | 15 | 5 |
| atsz–73a A | tlantic Source | Zone -83.5657 | 16.6196 | 77.36 | 15 | 17.94 |
| atsz–73b A | tlantic Source | Zone -83.6650 | 17.0429 | 77.36 | 15 | 5 |
| atsz–74a A | tlantic Source | Zone -82.7104 | 16.7695 | 82.35 | 15 | 17.94 |
| atsz–74b A | tlantic Source | Zone -82.7709 | 17.1995 | 82.35 | 15 | 5 |
| atsz–75a A | tlantic Source | Zone -81.7297 | 16,9003 | 79.86 | 15 | 17.94 |
| atsz–75b A | tlantic Source | Zone -81.8097 | 17.3274 | 79.86 | 15 | 5 |
| atsz-76a A | tlantic Source | Zone -80.9196 | 16.9495 | 82.95 | 15 | 17.94 |
| atsz–76b A | tlantic Source | Zone -80.9754 | 17.3801 | 82.95 | 15 | 5 |
| atsz-77a A | tlantic Source | Zone -79.8086 | 17.2357 | 67.95 | 15 | 17.94 |
| atsz-77b A | tlantic Source | Zone -79.9795 | 17.6378 | 67.95 | 15 | 5 |
| atsz-78a A | tlantic Source | Zone -79.0245 | 17.5415 | 73.61 | 15 | 17.94 |
| atsz–78b A | tlantic Source | Zone -79,1532 | 17.9577 | 73.61 | 15 | 5 |
| atsz-79a A | tlantic Source | Zone -78.4122 | 17.5689 | 94.07 | 15 | 17.94 |
| atsz-79b A | tlantic Source | Zone -78.3798 | 18 0017 | 94.07 | 15 | 5 |
| atsz-80a A | tlantic Source | Zone -77 6403 | 17 4391 | 103.3 | 15 | 17 94 |
| atsz–80b A | tlantic Source | Zone -77 5352 | 17.8613 | 103.3 | 15 | 5 |
| atsz-81a A | tlantic Source | Zone -76.6376 | 17.2984 | 98.21 | 15 | 17.94 |
| atsz–81b A | tlantic Source | Zone -76.5726 | 17 7278 | 98.21 | 15 | 5 |
| atsz-82a | tlantic Source | Zone -75 7299 | 19.0217 | 260.1 | 15 | 17 94 |
| atsz–82b A | tlantic Source | Zone -75.6516 | 18 5942 | 260.1 | 15 | 5 |
| atsz-83a A | tlantic Source | Zone -74 8351 | 19 2911 | 260.1 | 15 | 17 94 |
| atsz-83b A | tlantic Source | Zone -74 7621 | 18 8628 | 260.8 | 15 | 5 |
| atsz-84a A | tlantic Source | Zone -73.6639 | 19 2991 | 200.0 | 15 | 17 94 |
| atsz-84b A | tlantic Source | Zone -73 7026 | 18 8668 | 274.0 | 15 | 5 |
| atsz 040 f | tlantic Source | $Z_{one} = -72.8108$ | 10.0000 | 274.0 | 15 | 17.94 |
| atsz 85b A | tlantic Source | $Z_{\text{one}} = \frac{72.0130}{72.8246}$ | 18 7681 | 270.0 | 15 | 5 |
| atsz-600 A | tlantic Source | $Z_{\text{ODE}} = -72.8240$ $Z_{\text{ODE}} = 71.0143$ | 10.1031 | 270.0 | 15 | 17.04 |
| atsz-86b A | tlantic Source | Z_{0} -71.9145 Z_{0} -71.0068 | 19.1477 | 209.1 | 15 | 11.54 |
| atsz=800 A | tlantic Source | $Z_{\text{ODE}} = -71.9008$ | 18 8891 | 209.1 | 15 | 17.04 |
| atsz-07a A | tlantic Source | $Z_{\text{one}} = -70.4738$ | 18.5245 | 204.5 | 15 | 11.34 |
| atsz-orb A | tlantic Source | $Z_{one} = -70.7529$ | 18,0240 | 202.0 | 15 | 17.04 |
| atsz-ooa A | tlantic Source | $Z_{one} = -09.7710$ | 18.5902 | 200.8 | 15 | 11.94 |
| atsz-ood A | tlantic Source | $Z_{one} = -70.0347$ | 18.0004 | 306.4 | 15 | 17.04 |
| atoz-09a P | tlantic Source | Zone -03.2030 Zone 60.2790 | 10.2099 | ⊿໐ ວ .ສ ໑໑໑ ∩ | 10 | 11.94 5 |
| atoz 000 A | tlantic Source | 7_{ODO} 68_{EOEO} | 10 1449 | 200.8 979.0 | 10 | 17.04 |
| atsz-90a A | tlantic Source | Zone -00.0009 | 10.1443 17.7110 | 272.9 | 10 | 17.94 E |
| atsz-900 A | tlantic Source | $Z_{0110} = -00.0284$ | 10 1490 | 212.9 | 10 | 0 17.04 |
| atsz-91a A | tlantic Source | Zone -07.0428 | 18.1438 | 201.8 | 15 | 17.94 |
| atsz-910 P | tiantic Source | Zone -07.0200 | 10.0590 | 207.8 | 10 | 0 17.04 |
| atsz-92a A | tiantic Source | Zone -66.8261 | 18.2536 | 262 | 15 | 17.94 |
| atsz-920 A | tiantic Source | Lone -66.7627 | 17.8240 | 262 | 15 | Б |

Table B.1 – continued from previous page


Figure B.2: South Sandwich Islands Subduction Zone.

Table B.2: Earthquake parameters for South Sandwich Islands Subduction Zone unit sources.

| _ | Segment | Description | $Longitude(^{o}E)$ | $Latitude(^{o}N)$ | $\operatorname{Strike}(^{\mathrm{o}})$ | $\operatorname{Dip}(^{\mathrm{o}})$ | Depth (k | m) | |
|----------|---------|----------------|--------------------|-------------------|--|-------------------------------------|----------|-------|-------|
| sssz-1a | South | Sandwich Islan | ds Subduction Zor | e -32.3713 | -55.4 | 655 | 104.7 | 28.53 | 17.51 |
| sssz–1b | South | Sandwich Islan | ds Subduction Zor | e -32.1953 | -55.0 | 832 | 104.7 | 9.957 | 8.866 |
| sssz-1z | South | Sandwich Islan | ds Subduction Zor | e -32.5091 | -55.7 | 624 | 104.7 | 46.99 | 41.39 |
| sssz–2a | South | Sandwich Islan | ds Subduction Zor | e -30.8028 | -55.6 | 842 | 102.4 | 28.53 | 17.51 |
| sssz–2b | South | Sandwich Islan | ds Subduction Zor | e -30.6524 | -55.2 | 2982 | 102.4 | 9.957 | 8.866 |
| sssz-2z | South | Sandwich Islan | ds Subduction Zor | e -30.9206 | -55.9 | 839 | 102.4 | 46.99 | 41.39 |
| sssz-3a | South | Sandwich Islan | ds Subduction Zor | e -29.0824 | -55.8 | 3403 | 95.53 | 28.53 | 17.51 |
| sssz-3b | South | Sandwich Islan | ds Subduction Zor | e -29.0149 | -55.4 | 468 | 95.53 | 9.957 | 8.866 |
| sssz-3z | South | Sandwich Islan | ds Subduction Zor | e -29.1353 | -56.1 | 458 | 95.53 | 46.99 | 41.39 |
| sssz-4a | South | Sandwich Islan | ds Subduction Zor | e -27.8128 | -55.9 | 796 | 106.1 | 28.53 | 17.51 |
| sssz-4b | South | Sandwich Islan | ds Subduction Zor | e -27.6174 | -55.5 | 999 | 106.1 | 9.957 | 8.866 |
| sssz-4z | South | Sandwich Islan | ds Subduction Zor | e -27.9659 | -56.2 | 2744 | 106.1 | 46.99 | 41.39 |
| sssz-5a | South | Sandwich Islan | ds Subduction Zor | e -26.7928 | -56.2 | 2481 | 123.1 | 28.53 | 17.51 |
| sssz-5b | South | Sandwich Islan | ds Subduction Zor | e -26.4059 | -55.9 | 170 | 123.1 | 9.957 | 8.866 |
| sssz-5z | South | Sandwich Islan | ds Subduction Zor | e -27.0955 | -56.5 | 052 | 123.1 | 46.99 | 41.39 |
| sssz-6a | South | Sandwich Islan | ds Subduction Zor | e -26.1317 | -56.6 | 5466 | 145.6 | 23.28 | 16.11 |
| sssz-6b | South | Sandwich Islan | ds Subduction Zor | e -25.5131 | -56.4 | 133 | 145.6 | 9.09 | 8.228 |
| sssz-6z | South | Sandwich Islan | ds Subduction Zor | e -26.5920 | -56.8 | 3194 | 145.6 | 47.15 | 35.87 |
| sssz-7a | South | Sandwich Islan | ds Subduction Zor | e -25.6787 | -57.2 | 2162 | 162.9 | 21.21 | 14.23 |
| sssz-7b | South | Sandwich Islan | ds Subduction Zor | e -24.9394 | -57.0 | 932 | 162.9 | 7.596 | 7.626 |
| sssz-7z | South | Sandwich Islan | ds Subduction Zor | e -26.2493 | -57.3 | 3109 | 162.9 | 44.16 | 32.32 |
| sssz-8a | South | Sandwich Islan | ds Subduction Zor | e -25.5161 | -57.8 | 3712 | 178.2 | 20.33 | 15.91 |
| sssz-8b | South | Sandwich Islan | ds Subduction Zor | e -24.7233 | -57.8 | 3580 | 178.2 | 8.449 | 8.562 |
| sssz-8z | South | Sandwich Islan | ds Subduction Zor | -26.1280 | -57.8 | 813 | 178.2 | 43.65 | 33.28 |
| sssz-9a | South | Sandwich Islan | ds Subduction Zor | e -25.6657 | -58.5 | 0053 | 195.4 | 25.76 | 15.71 |
| sssz-9b | South | Sandwich Islan | ds Subduction Zor | e -24.9168 | -58.6 | 5127 | 195.4 | 8.254 | 8.537 |
| sssz-9z | South | Sandwich Islan | ds Subduction Zor | e -26.1799 | -58.4 | 313 | 195.4 | 51.69 | 37.44 |
| sssz-10a | South | Sandwich Islan | ds Subduction Zor | e -26.1563 | -59.1 | .048 | 212.5 | 32.82 | 15.65 |
| sssz-10b | South | Sandwich Islan | ds Subduction Zor | e -25.5335 | -59.3 | 080 | 212.5 | 10.45 | 6.581 |
| sssz-10z | South | Sandwich Islan | ds Subduction Zor | e -26.5817 | -58.9 | 653 | 212.5 | 54.77 | 42.75 |
| sssz-11a | South | Sandwich Islan | ds Subduction Zor | e -27.0794 | -59.6 | 5799 | 224.2 | 33.67 | 15.75 |
| sssz-11b | South | Sandwich Islan | ds Subduction Zor | e -26.5460 | -59.9 | 412 | 224.2 | 11.32 | 5.927 |
| sssz–11z | South | Sandwich Islan | ds Subduction Zor | e -27.4245 | -59.5 | 6098 | 224.2 | 57.19 | 43.46 |

Appendix C SIFT Testing Report

Appendix C

SIFT Testing Report

Authors: Lindsey Wright, Liujuan Tang

1.0 PURPOSE

Forecast models are tested with synthetic tsunami events covering a range of tsunami source locations and magnitudes. Testing is also done with selected historical tsunami events when available.

The testing of a forecast model has three objectives. The first objective is to assure that the results obtained with the NOAA's tsunami forecast system software, which has been released to the Tsunami Warning Centers for operational use are consistent with those obtained by the researcher during the development of the forecast model. The second objective is to test the forecast model for consistency, accuracy, time efficiency, and quality of results over a range of possible tsunami locations and magnitudes. The third objective is to identify bugs and issues in need of resolution by the researcher who developed the Forecast Model or by the forecast system software development team before the next version release to NOAA's two Tsunami Warning Centers.

Local hardware and software applications, and tools familiar to the researcher(s), are used to run the Method of Splitting Tsunamis (MOST) model during the forecast model development. The test results presented in this report lend confidence that the model performs as developed and produces the same results when initiated within the forecast system application in an operational setting as those produced by the researcher during the forecast model development. The test results assure those who rely on the Virginia Beach tsunami forecast model that consistent results are produced irrespective of system.

2.0 TESTING PROCEDURE

The general procedure for forecast model testing is to run a set of synthetic tsunami scenarios and a selected set of historical tsunami events through the forecast system application and compare the results with those obtained by the researcher during the forecast model development and presented in the Tsunami Forecast Model Report. Specific steps taken to test the model include:

- 1. Identification of testing scenarios, including the standard set of synthetic events, appropriate historical events, and customized synthetic scenarios that may have been used by the researcher(s) in developing the forecast model.
- 2. Creation of new events to represent customized synthetic scenarios used by the researcher(s) in developing the forecast model, if any.
- 3. Submission of test model runs with the forecast system, and export of the results from A, B, and C grids, along with time series.
- 4. Recording applicable metadata, including the specific forecast system version used for testing.
- 5. Examination of forecast model results for instabilities in both time series and plot results.
- 6. Comparison of forecast model results obtained through the forecast system with those obtained during the forecast model development.
- 7. Summarization of results with specific mention of quality, consistency, and time efficiency.
- 8. Reporting of issues identified to modeler and forecast system software development team.
- 9. Retesting the forecast models in the forecast system when reported issues have been addressed or explained.

Synthetic model runs were tested on a DELL PowerEdge R510 computer equipped with two Xeon E5670 processors at 2.93 Ghz, each with 12 MBytes of cache and 32GB memory. The processors are hex core and support hyperthreading, resulting in the computer performing as a 24 processor core machine. Additionally, the testing computer supports 10 Gigabit Ethernet for fast network connections. This computer configuration is similar or the same as the configurations of the computers installed at the Tsunami Warning Centers so the compute times should only vary slightly.

Results

The Virginia Beach forecast model was tested with SIFT version 3.2. The same version of propagation database was used during the model development.

The Virginia Beach, Virginia forecast model was tested with three synthetic scenarios. Test results from the forecast system and comparisons with the results obtained during the forecast model development are shown numerically in Table C1 and graphically in Figures C1 to C3. The results show that the minimum and maximum amplitudes and time series obtained from the forecast system agree with those obtained during the forecast model development, and that the forecast model is stable and robust, with consistent and high quality results across geographically distributed tsunami sources. The model run time (wall clock time) was 24.75 minutes for 11.99 hours of simulation time, and 8.24 minutes for 4.0 hours. This run time is well within the 10 minute run time for 4 hours of simulation time.

A suite of three synthetic events was run on the Virginia Beach forecast model. The modeled scenarios were stable for all cases run. The largest modeled height was 601 centimeters (cm) from the Atlantic (ATSZ 48–57) source zone. The smallest signal of 40 cm was recorded at the South Sandwich (SSSZ 1–10) source zone. Maximum values for the SSSZ 1–10 differed slightly. This abnormality resulted from the development model being run for a longer window of time than the SIFT output and therefore an additional wave with a slightly higher amplitude was recorded at approximately 28 hours after the event. Visual comparisons between the development cases and the forecast system output were nearly identical in shape and amplitude for all cases. The Virginia Beach reference point used for the forecast model development is the same as what is deployed in the forecast system, so the results can be considered valid for the three cases studied.

 Table C1. Table of maximum and minimum amplitudes (cm) at the Virginia Beach, Virginia warning point for synthetic and historical events

 tested using SIFT 3.2 and obtained during development.

| Scenario Name | Source Zone | Tsunami Source | α [m] | SIFT Max | Developme | SIFT Min | Developmen t Min (cm) | | | | |
|------------------------|----------------|------------------|----------|----------|--------------|----------|--------------------------|--|--|--|--|
| Hume | | | [] | (em) | ine max (em) | (em) | c Mill (cill) | | | | |
| Mega-tsunami Scenarios | | | | | | | | | | | |
| ATSZ 38-47 | Atlantic | A38-A47, B38-B47 | 25 | 246.9 | 247 | -122.8 | -123 | | | | |
| ATSZ 48-57 | Atlantic | A48-A57, B48-B57 | 25 | 600.9 | 601 | -356.9 | -358 | | | | |
| SSSZ 1-10 | South Sandwich | A1-A10, B1-B10 | 25 | 39.9 | 58(40 for | -40.6 | | | | | |
| | | | | | first 12 | | -43 | | | | |
| | | | | | hours) | | | | | | |



Figure C1 Response of the Virginia Beach forecast model to synthetic scenario ATSZ 38-47. (a, b, and c) Maximum sea surface elevation for A-, B- and C-grids. (d) Sea surface elevation time series at the C-grid warning point.



Figure C2 Response of the Virginia Beach forecast model to synthetic scenario ATSZ 48-57. (a, b, and c) Maximum sea surface elevation for A-, B- and C-grids. (d) Sea surface elevation time series at the C-grid warning point.



Figure C3 Response of the Virginia Beach forecast model to synthetic scenario SSSZ 1-10. (a, b, and c) Maximum sea surface elevation for A-, B- and C-grids. (d) Sea surface elevation time series at the C-grid warning point.